Resonant excitation of trapped coastal waves by free inertia-gravity waves and its effects on transport and mixing

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Ideology

- Trapped waves in the coastal waveguide interact with free inertia-gravity waves
- Cubic nonlinearity \rightarrow interaction triadic
- Interaction may be resonant ↔ resonant triads
- If so, resonant excitation of trapped waves by free waves, with two possible mechanisms:
 - Free wave \rightarrow Trapped wave + Trapped wave ($F \rightarrow TT$)
 - 3 Free wave + Free wave \rightarrow Trapped wave ($F + F \rightarrow T$)
- Resonant growth of the trapped wave + nonlinearity/dissipation ⇒ coherent/dissipative structures

Remark

Twist: infinitely long trapped wave \equiv Coastal current \Rightarrow *F* + *C* \rightarrow *T* resonance.

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Setup: one- or two-layer rotating shallow water model



Steep vs gentle bathymetry

Steep slopes

Characteristic horizontal scale of bathymetry $\langle R_d \Rightarrow$ Trapped waves (leading order): (almost) non-dispersive Kelvin waves. Possible resonances:

 $F + F \rightarrow T, F + C \rightarrow T$ (baroclinic case).

Gentle slope

Characteristic horizontal scale of bathymetry $\mathcal{O}(R_d) \Rightarrow$ Trapped waves: dispersive shelf and edge waves. Details of bathymetry unimportant: spectrum of coastal waves universal, Huthnance, 1975. All resonances possible.

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Barotropic 1-layer model Baroclinic 2-layer model

Non-dimensional RSW equations with idealized coast

$$u_t - v + h_x = -\epsilon(uu_x + vu_y)$$

$$v_t + u + h_y = -\epsilon(uv_x + vv_y)$$

$$h_t + u_x + v_y = -\epsilon((hu)_x + (hv)_y).$$

Boundary condition: $x \ge 0$, $u|_{x=0} = 0$.

Scaling

Time:
$$f^{-1}$$
; Space: $L \sim R_d = \frac{\sqrt{gH}}{f}$; Velocity: *U*.
Rossby number $Ro \equiv \epsilon = \frac{U}{fL}$.

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Linear waves $(u, v, h) \propto e^{i(\sigma t - kx - ly)}$:

• Free inertia-gravity waves (IGW) with dispersion

$$\sigma = \sqrt{1 + k^2 + l^2}$$

• Trapped non-dispersive Kelvin waves (KW) with

$$\sigma = -l, \quad k = -i$$

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Ideology and setup Steep slope Conclusions

Barotropic 1-layer model

IGW - KW interaction



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 Coastal waves: Resonant Excitation, Transport and Mixing

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Conditions of IGW - KW resonance

A pair of IGW with frequencies $\sigma_{1,2}$ and along-coast wavenumbers $I_{1,2}$ is in resonance with a KW with wavenumber Iif ("difference" resonance)

$$\sigma_1 - \sigma_2 = \sigma_K = -I, \quad I_1 - I_2 = I, \ I \neq 0.$$
 (1)

For *I* < 0:

$$|I| = \sqrt{1 + k_1^2 + l_1^2} - \sqrt{1 + k_2^2 + l_2^2}, \quad l_2 = l_1 + |I|, \quad (2)$$

and

$$\sqrt{1+k_1^2+l_1^2}-|l|=\sqrt{1+k_2^2+(l_1+|l|)^2}.$$
 (3)

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Evolution equation for KW

Killing resonances in RSW equations with coast with the help of slow-time corrections \Rightarrow

$$K_{T} + KK_{\eta} = Se^{il\eta} + S^{*}e^{-il\eta}, \qquad (4)$$

where $\eta = y + t$,

$$S = \int_{0}^{\infty} dx \, e^{-x} \left[(H_{1} U_{2}^{*} + U_{1} H_{2}^{*})_{x} - U_{1} V_{2x}^{*} - V_{1x} U_{2} + i I (H_{1} V_{2}^{*} + V_{1} H_{2}^{*} - V_{1} V_{2}^{*}) \right], \qquad (5)$$

and (U_i, V_i, H_i) , i = 1, 2 are amplitudes of 2 free waves.

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Final form of the evolution equation

From the polarisation relations get:

$$S = iA_1A_2s, \quad Im(s) = 0 \tag{6}$$

and hence

$$K_T + KK_\eta = -2sA_1A_2\sin l\eta, \qquad (7)$$

- harmonically forced Hopf equation

$$s = \frac{4l}{(k_1^2 + 1)(k_2^2 + 1)[1 + (k_1 + k_2)^2][1 + (k_1 - k_2)^2]} \times \left[(\sigma_1 l_2 + \sigma_2 l_1 - l_1 l_2)(1 + k_1^2 + k_2^2) + \sigma_2 l_1 k_1 (1 + k_1^2 - k_2^2) + \sigma_1 l_2 k_2 (1 + k_2^2 - k_1^2) + 2k_1 k_2 (l_1 l_2 - (1 + k_1^2)(1 + k_2^2)) \right]$$
(8)

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Isopleths of the interaction coefficient $s(l, k_1, l_1)$ for l = -1 at the interval 10



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Comments

- Equation (7) is completely integrable in Lagrangian coordinates. Main property of solutions: breaking and formation of Kelvin fronts → a route to dissipation in the ocean
- The mechanism produces coastal wave "from nothing"
- Small deviations from infinite slope weak dispersion ⇒ forced Hopf equation → forced KdV equation, also integrable
- Details: Reznik & Zeitlin, Phys. Letters A, 2009.

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2-layer RSW with idealized coast



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Non-dimensional equatins of motion

$$\begin{pmatrix} \partial_t + \epsilon \vec{\mathbf{v}}_1 \cdot \vec{\nabla} \end{pmatrix} \vec{\mathbf{v}}_1 + f\hat{\mathbf{z}} \wedge \vec{\mathbf{v}}_1 + \vec{\nabla} (\eta_1 + \eta_2) &= \mathbf{0}, \\ \partial_t \eta_1 + \vec{\nabla} \cdot ((\mathbf{d} + \epsilon \eta_1) \vec{\mathbf{v}}_1) &= \mathbf{0}, \\ \left(\partial_t + \epsilon \vec{\mathbf{v}}_1 \cdot \vec{\nabla} \right) \vec{\mathbf{v}}_2 + f\hat{\mathbf{z}} \wedge \vec{\mathbf{v}}_2 + \vec{\nabla} (r\eta_1 + \eta_2) &= \mathbf{0}, \\ \partial_t \eta_2 + \vec{\nabla} \cdot ((1 - \mathbf{d} + \epsilon \eta_2) \vec{\mathbf{v}}_2) &= \mathbf{0}.$$

$$(9)$$

Scaling: same as in 1-layer RSW. New parameters: $d = \frac{H_1}{H_1 + H_2}$, $r = \frac{\rho_1}{\rho_2} \le 1$.

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Baroclinic-barotropic decomposition and wave spectrum

$$d = \frac{1}{2} \rightarrow \vec{v}^{\pm} = \sqrt{r} \vec{v}_1 \pm \vec{v}_2, \quad \eta^{\pm} = 2 \left(\sqrt{r} \eta_1 \pm \eta_2 \right) \Rightarrow$$

Wave spectrum

• Free IGW with dispersion relation:

$$\sigma_{IG^{\pm}}^{2} = 1 + c_{\pm}^{2} \vec{k}^{2}, \ \vec{k} = (k, I),$$
(10)

 Trapped dispersionless Kelvin (K) waves propagating along the y axis with the phase velocity c_±:

$$\sigma_{K^{\pm}}^2 = c_{\pm} l^2. \tag{11}$$

 $c_{\pm}=\sqrt{\frac{1\pm\sqrt{r}}{2}}$ - dimensionless phase velocities of BT and BC GW

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Wave and mean-current solutions

Waves

$$(u_{IG}^{\pm}, v_{IG}^{\pm}, \eta_{IG}^{\pm}) = (U^{\pm}(x), V^{\pm}(x), N^{\pm}(x)) e^{i(ly - \sigma t)} + c.c., \quad (12)$$
$$(u_{K}^{\pm}, v_{K}^{\pm}, \eta_{K}^{\pm}) = (0, c_{\pm}K^{\pm}(y + c_{\pm}t), -K^{\pm}(y + c_{\pm}t))e^{-x/c_{\pm}}, \quad (13)$$

 ${\cal K}^{\pm}$ - arbitrary functions of arguments $\xi^{\pm} = y + c_{\pm} t$

Mean current

$$u_M^{\pm} = 0, \quad v_M^{\pm} = 1/2(1 \pm \sqrt{r})\partial_x \eta_M^{\pm}.$$
 (14)

Coastal current: $v_{1,2}$ rapidly decaying. Surface current:

$$v_{2_M} = 0, \Rightarrow r\eta_{1_M} + \eta_{2_M} = 0.$$
 (15)

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Dispersion relation for 2-layer RSW in a half-plane

Wave spectrum: barotropic (BT) (faster) and baroclinic (BC) (slower) versions of each kind of waves: IGW and KW (baroclinic KW not shown).



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What is new due to baroclinicity?

• IGW+IGW \rightarrow KW resonances:

$$I_{IG_1} \pm I_{IG_2} = I_K, \quad \sigma_{IG}(\vec{k}_1) \pm \sigma_{IG}(\vec{k}_2) = \sigma_K(I),$$

- BT KW may be excited by BT BT, BT + BC, and BC + BC IGW ("difference" and "sum" resonances)
- BC KW may be excited by BT BT, BT BC, and BC BC IGW (only "difference" resonances)
- New resonance: BC IGW + Coastal Current (CC) → BT KW due to the intersection of dispersion surfaces

Remark

BC + BC IGW and BC IGW + CC excitation: weak wave signature at the surface \Rightarrow strong surface signature of the excited wave ("wave from nothing").

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Analysis of resonance conditions: wave-wave

"Difference" resonance conditions: similar to 1-layer case "Sum" resonance conditions:

$$I_{IG_1} + I_{IG_2} = I_K, \quad \sigma_{IG}(\vec{k}_1) + \sigma_{IG}(\vec{k}_2) = \sigma_K(I) \Rightarrow$$
(16)

$$l_{1}+l_{2} = l = -|l|, \quad \sqrt{1 + c_{-}^{2}(k_{1}^{2} + l_{1}^{2})} + \sqrt{1 + c_{-}^{2}(k_{2}^{2} + l_{2}^{2})} = c_{+}|l|, \quad (17)$$

Renormalizing the wavevectors $(k_i, l_i) = c_-^{-1}(\tilde{k}_i, \tilde{l}_i)$ gives

$$\sqrt{1+k_1^2+l_1^2}+\sqrt{1+k_2^2+l_2^2}=-\frac{c_+}{c_-}(l_1+l_2).$$
 (18)

Condition of existence of resonance for negative $I_{1,2}$:

$$\frac{c_{+}}{c_{-}}\left(|l|_{1}+|l|_{2}\right) > \sqrt{1+|l|_{1}^{2}} + \sqrt{1+|l|_{2}^{2}}$$
(19)

- satisfied for weak enough stratifications.

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Analysis of resonance conditions: wave-mean

Resonance condition:

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$$I_{G^{-}} = \sigma_{K^{+}}, \quad I_{IG^{-}} = I_{K^{+}} \equiv I \Rightarrow$$

 $c_{+}^{2}I^{2} = 1 + c_{-}^{2}(k^{2} + l^{2}),$ (20)

which leads to the following expression for the x- component k of the wavenumber of the incoming IG wave:

$$k^2 = \left(\frac{c_+^2}{c_-^2} - 1\right) l^2 - \frac{1}{c_-^2}.$$
 (21)

For given stratification, i.e. $\frac{c_{+}^{2}}{c_{-}^{2}} > 1$, the resonance is always possible for sufficiently large negative *I*.

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Evolution equation for the Kelvin wave envelopes

Wave-wave excitation

$$\mathcal{K}_{\mathcal{T}}^{\pm} + \mathbf{a}_{\pm}\mathbf{c}_{\pm}\mathcal{K}^{\pm}\mathcal{K}_{\xi^{\pm}}^{\pm} = \mathcal{S}, \tag{22}$$

where ${\mathcal S}$ - forcing by nonlinear interaction of IGW, harmonic in $\xi^\pm.$

Wave-mean excitation

$$K_T^+ + CK_{\xi^+}^+ + a_+c_+K^+K_{\xi^+}^+ = S,$$
 (23)

C- velocity induced by the mean current (Doppler shift); S proportional to the amplitude of the IGW and harmonic in ξ^+

Here

$$a_{\pm} = \frac{1}{4} \left(\frac{1}{\sqrt{r}} \pm 1 \right), \qquad (24)$$

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RSW equations with coast

Equations of motion:

$$u_{t} + uu_{x} + vu_{y} - fv + g\eta_{x} = 0,$$

$$v_{t} + uv_{x} + vv_{y} + fu + g\eta_{y} = 0,$$

$$\eta_{t} + ((\eta + h)u)_{x} + ((\eta + h)v)_{y} = 0.$$
 (25)

Water depth h(x) - 1d topography. Coast:

$$\left. h \right|_{x=0} = 0, \quad \left. h \right|_{x \to \infty} \to H = \text{const.}$$
 (26)

Boundary conditions:

$$\eta|_{x \to 0}$$
 regular, $\eta u|_{x \to 0} \to 0.$ (27)

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Linearized system

Linearized non-dimensional equations:

$$u_{t} - fv + \eta_{x} = 0,$$

$$v_{t} + fu + \eta_{y} = 0,$$

$$\eta_{t} + (hu)_{x} + (hv)_{y} = 0,$$
(28)

• Trapped waves:

$$(\eta, \boldsymbol{u}, \boldsymbol{v})|_{\boldsymbol{x} \to \infty} \to \boldsymbol{0}.$$
 (29)

• Free (incident + reflected) wave:

$$(\eta, u, v)|_{x \to \infty} \propto e^{i(ly - \sigma t)} Re \left(A^+ e^{ikx} + A^- e^{-ikx} \right),$$
 (30)

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Linear waves

Wave equation:

Solutions of linearized system:

$$(u, v, \eta) = (iU, V, Z)(x)e^{i(ly - \sigma t)}, \Rightarrow$$
(31)

$$U = \frac{f I Z - \sigma Z'}{\sigma^2 - f^2}, \quad V = \frac{\sigma I Z - f Z'}{\sigma^2 - f^2}, \quad ' \equiv \frac{d}{dx}.$$
 (32)

Hence

$$\left(hZ'\right)' + \left(\sigma^2 - f^2 - l^2h - \frac{fl}{\sigma}h'\right)Z = 0, \qquad (33)$$

an eigenproblem for eigenfunctions Z_n and eigenfrequencies σ_n . Solution below: Ball's model $h = 1 - e^{-x}$ (Ball, 1968); similar for any monotonous profile (Huthnance, 1975).

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Linear wave spectrum of the RSW model with a gentle coastal slop

Dispersion diagram



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Weakly nonlinear system

Nonlinear equations:

$$u_t - fv + \eta_x = -\epsilon (uu_x + vu_y),$$

$$v_t + fu + \eta_y = -\epsilon (uv_x + vv_y),$$

$$\eta_t + (hu)_x + (hv)_y = -\epsilon ((\eta u)_x + (\eta v)_y),$$
 (34)

$$\epsilon = \frac{U}{\sqrt{gH}} << 1 - \text{Froude number.}$$

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Resonant triads

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• lowest-order solution (u_0, v_0, η_0) is in the form:

$$\mathbf{W}_0 + \mathbf{W}_1 + \mathbf{W}_2,$$

where $W_i = A_i(iU_i, V_i, Z_i)e^{i\theta_i}$, $\theta_i = I_iy - \sigma_i t$, i = 0, 1, 2 - a pair of trapped waves, i = 1, 2, and an incident/reflected wave i = 0,

and

• synchronism conditions hold:

$$\mathbf{I}_0 = \mathbf{I}_1 + \mathbf{I}_2, \quad \sigma_0 = \sigma_1 + \sigma_2.$$
 (35)

⇒ the amplitudes $A_{1,2}$ grow exponentially in slow time $T_2 = e^{-1}t.$

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Graphical solution for resonances



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Modulation equations

Coupled generalized Landau equations:

$$\partial_{\mathbf{T}_2} \mathbf{A}_1 + i\mathbf{M}_{11} |\mathbf{A}_1|^2 \mathbf{A}_1 + i\mathbf{M}_{12} |\mathbf{A}_2|^2 \mathbf{A}_1 + i\mathbf{M}_{02} \mathbf{A}_0 \mathbf{A}_2^* = 0$$

$$\partial_{\textbf{T}_2}\textbf{A}_2 + i\textbf{M}_{21} \left|\textbf{A}_1\right|^2 \textbf{A}_2 + i\textbf{M}_{22} \left|\textbf{A}_2\right|^2 \textbf{A}_2 + i\textbf{M}_{01}\textbf{A}_0\textbf{A}_1^* \hspace{2mm} = \hspace{2mm} 0$$

 M_{ij} - convolutions of zeroth and first orders for waves i and j, $M_{02}M_{01}^{\ast}>0$ (condition for resonant growth).

General property:

Saturation \equiv existence of attracting limit cycle, or attracting fixed point in A_1, A_2 - space (Details depend on the parameters of trapped waves).

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Comments

- If spatial modulation taken into account structure formation
- Earlier results (Miles, 1990) corrected and generalized for arbitrary incidence angle
- Details: Reznik & Zeitlin, JFM, 2011.

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Conclusions

Several mechanisms of resonant excitation of trapped coastal waves exist, either leading to formation of coherent/dissipative structures in the coastal zones. Resonant pairs of free inertia-gravity waves, or resonances of free inertia-gravity waves with coastal currents produce coastal waves "from nothing"

Important for:

- Energy and momentum transport from the open ocean to the coast, and subsequent dissipation. A poorly explored route to dissipation in the ocean.
- Transport and mixing in the coastal zones.