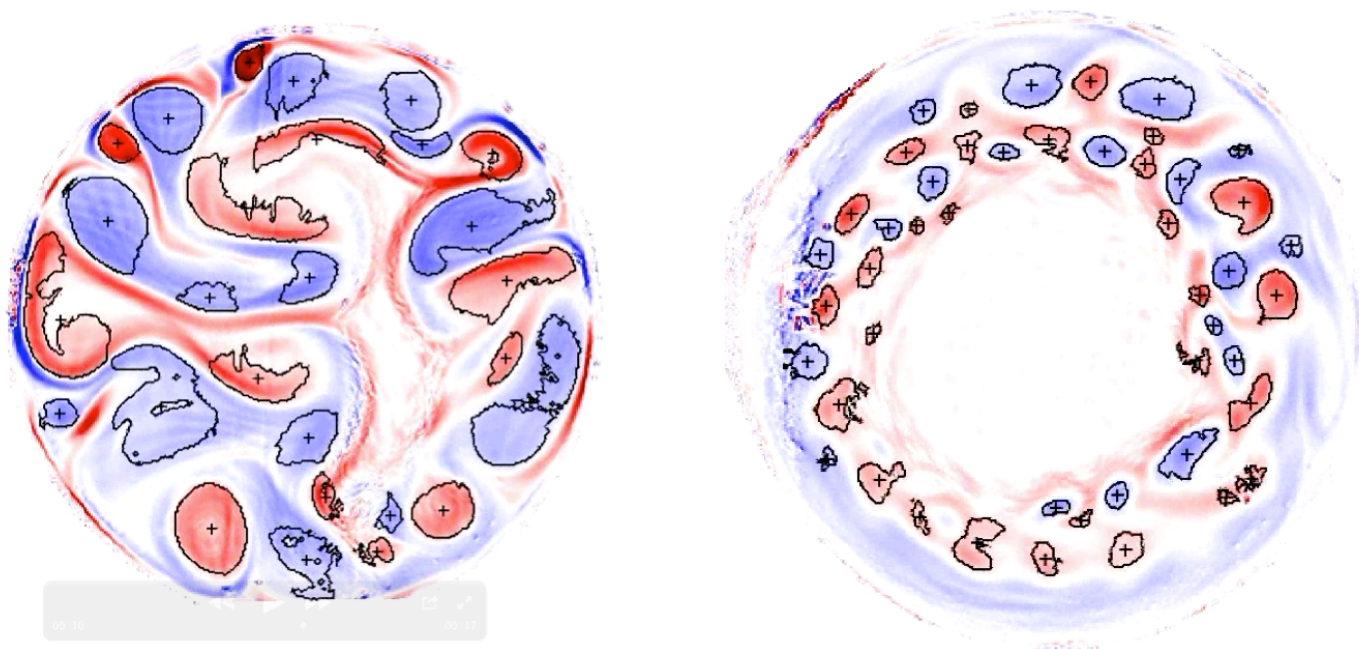


*Meso and sub meso scale dynamics
of coastal current along steep shelf bathymetry*



N.GEHENIAU⁽¹⁾ R. PENNEL^(1,2)

K. BERANGER⁽³⁾ F.POULIN⁽⁴⁾ A. STEGNER⁽¹⁾

⁽¹⁾ *Laboratoire de Météorologie Dynamique, CNRS, Ecole Polytechnique, Palaiseau, France.*

⁽²⁾ *Laboratoire de Physique des Océans, UBO, Brest, France.*

⁽³⁾ *UME, ENSTA-ParisTech, Palaiseau, France.*

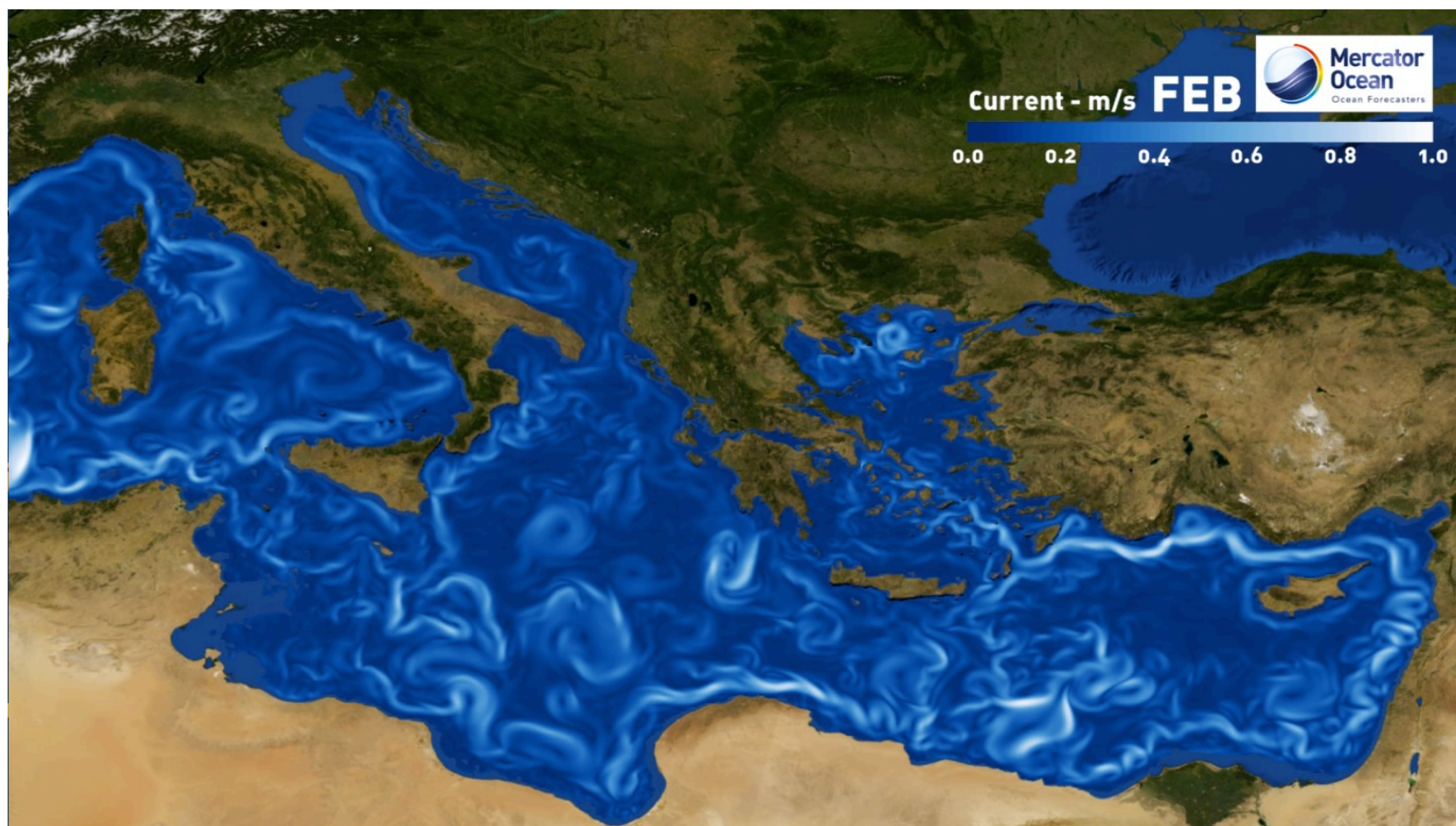
⁽⁴⁾ *Waterloo University, Canada.*



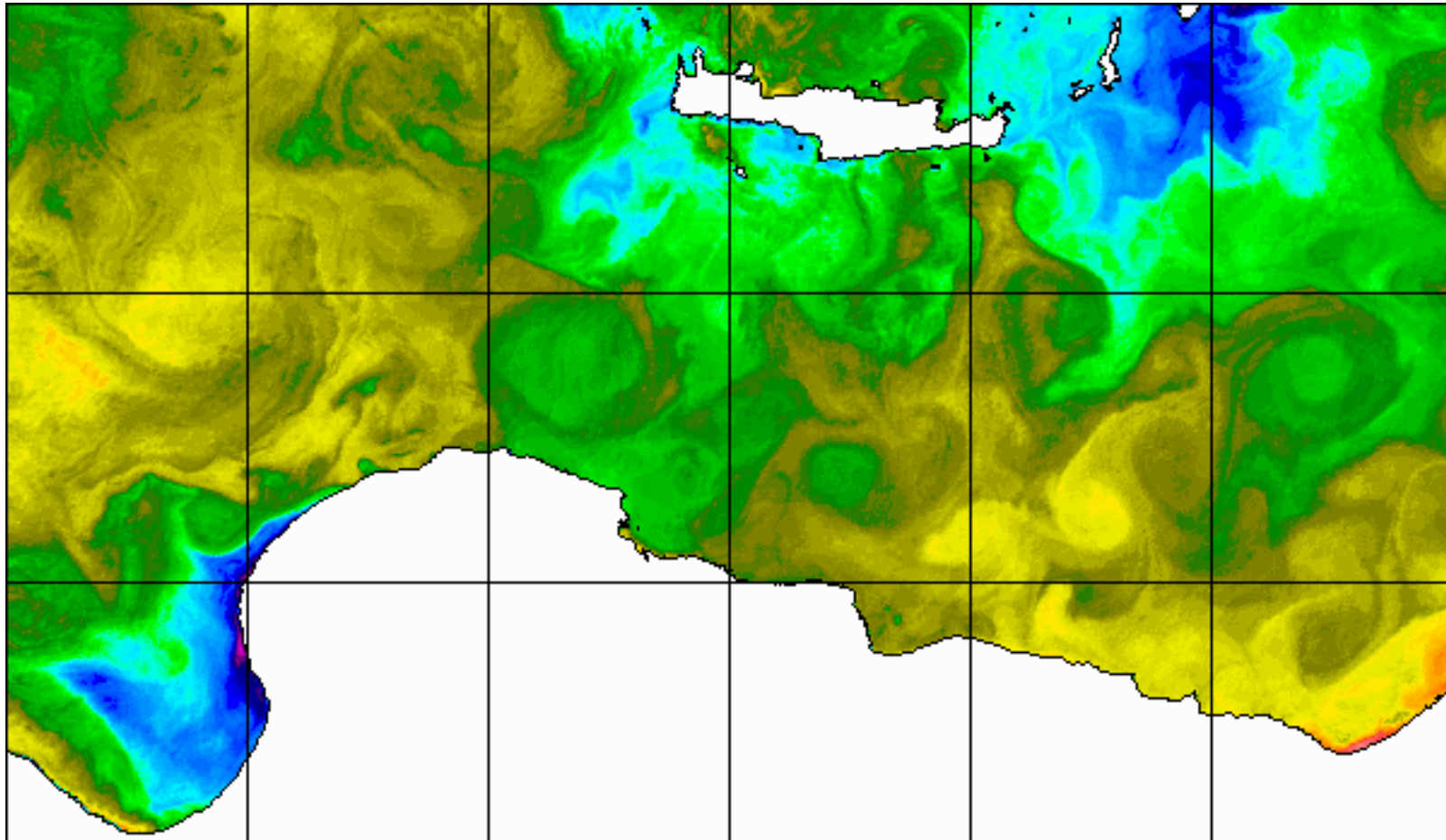
Coastal current along steep shelf bathymetry



Motivations: $1/36^\circ$ Eastern Mediterranean basin

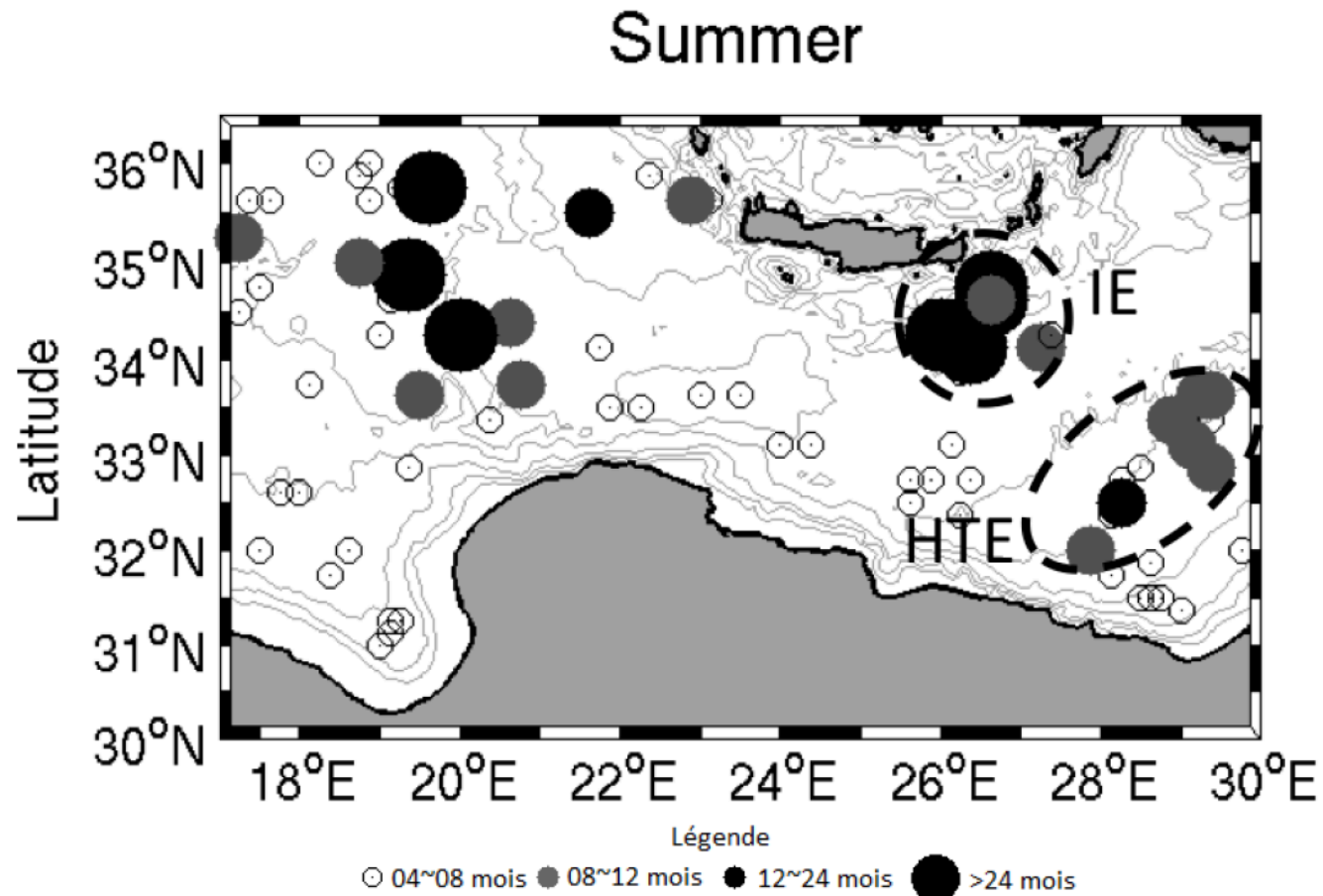


Motivations: formation of coastal eddies Mediterranean Sea



SST 18 June 2006

Motivations: formation area over steep or smooth shelf slope ?





Coastal current along steep shelf bathymetry

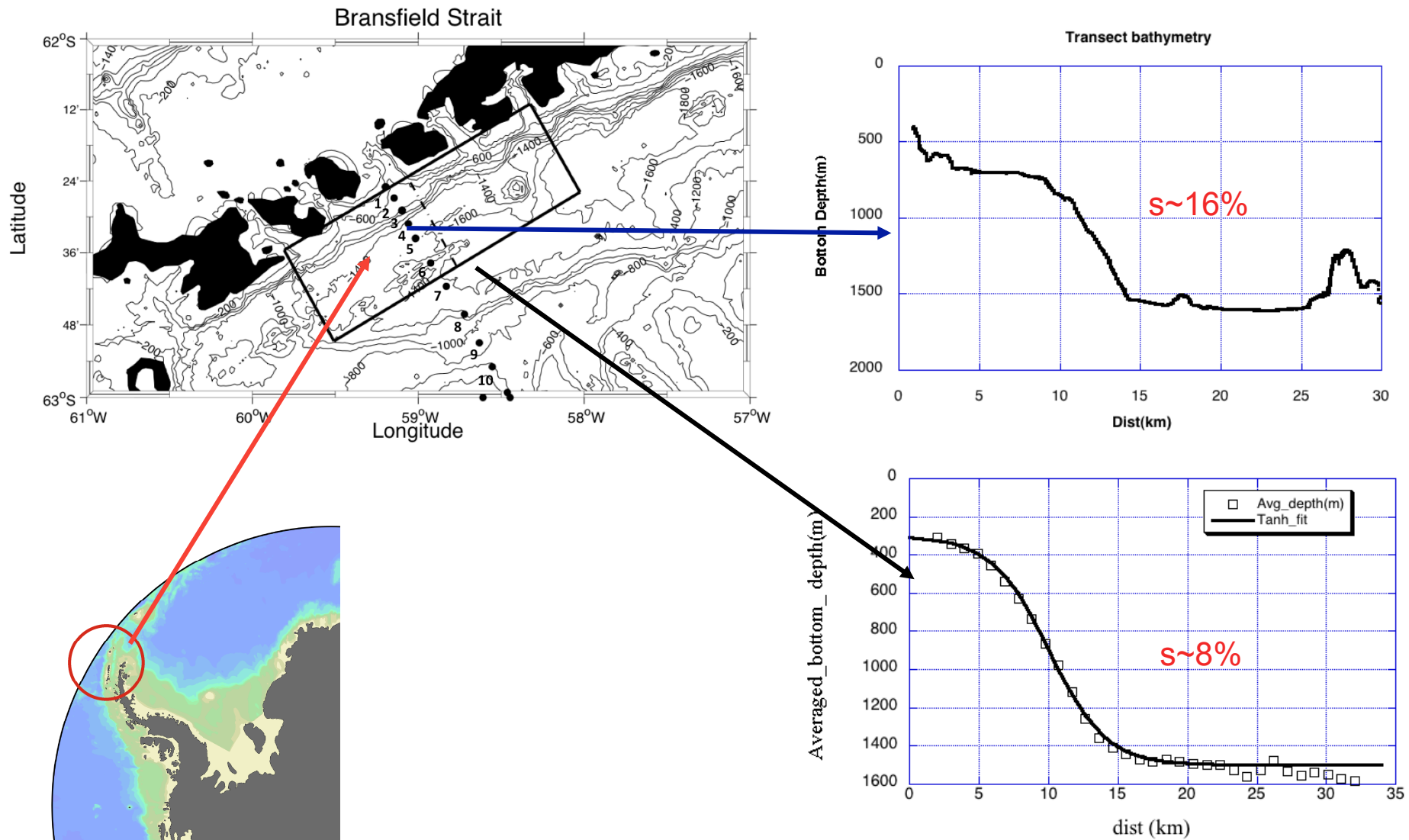


Motivations: Bransfield Current (Antarctica)



COUPLING in-situ survey 2010

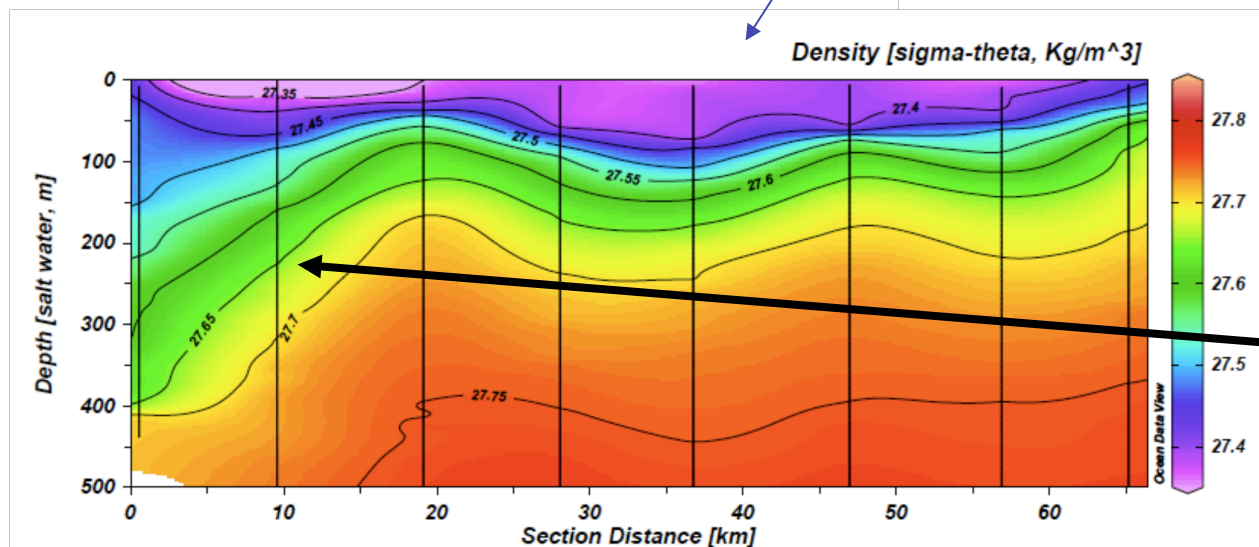
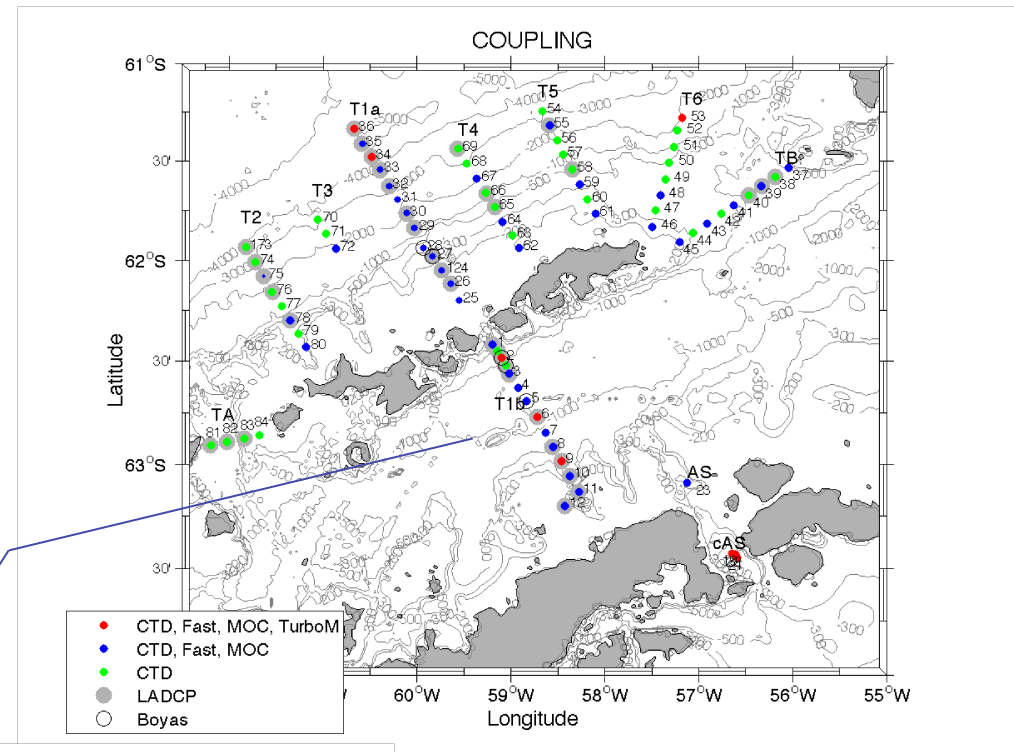
Buoyant coastal current : Bransfield strait bathymetry (Antarctica)



Buoyant coastal current :
Bransfield current (Antarctica)

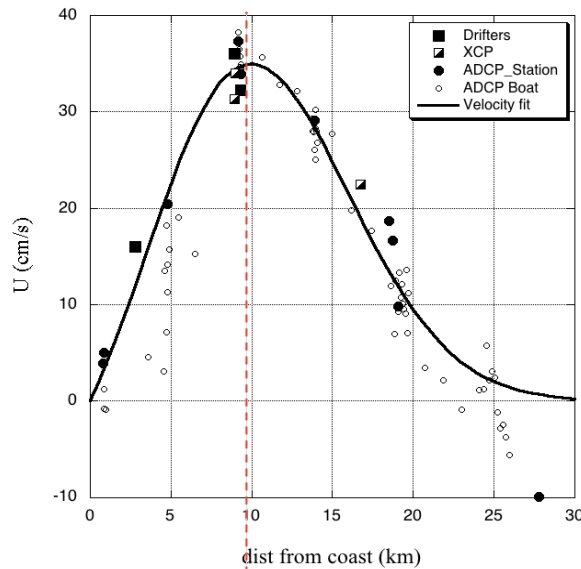
COUPLING Campaign 2010

CTD Transect 1b



Buoyant coastal current : stable Bransfield current (Antarctica)

VELOCITY PROFILE

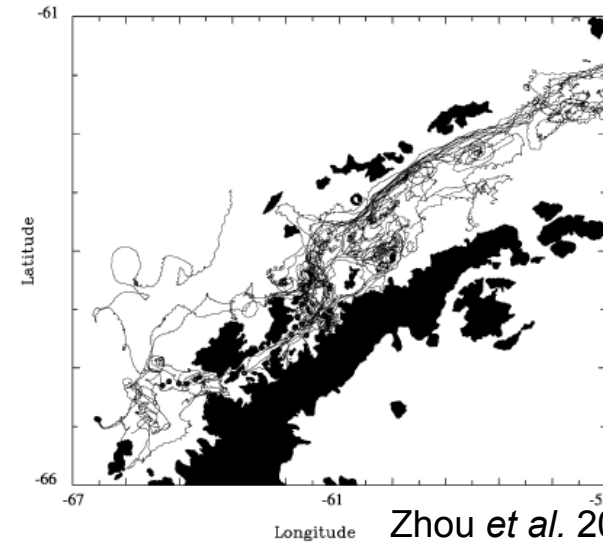


$$Ro = \frac{U}{fL} \approx 0.2 - 0.3$$

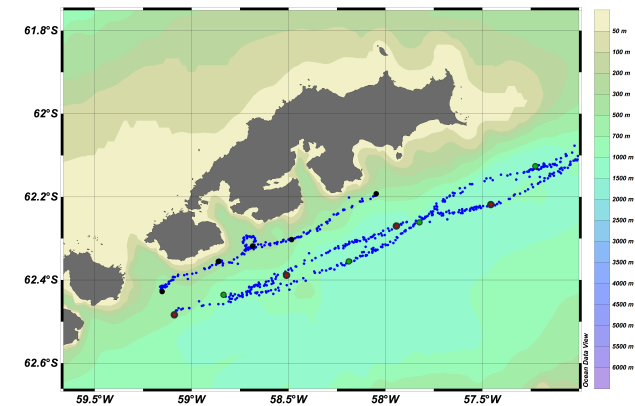
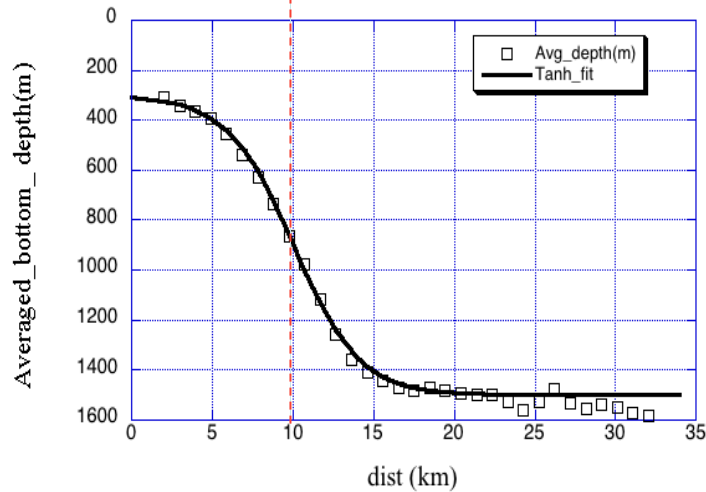
$$Bu = \left(\frac{R_d}{L} \right)^2 \approx 1$$

GEOSTROPHIC CURRENT

DRIFTERS TRAJECTORIES

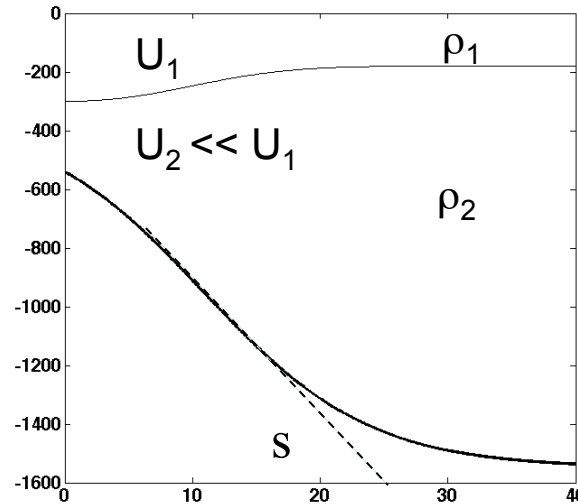


Zhou et al. 2002



Sangra et al. DSR 2011, Poulin et al. JPO 2014

Simple configuration



Surface intensified current

$$Ro \quad Bu = (R_d/L)^2$$

Vertical stratification

$$\gamma = H_1/H_2$$

Shelf bathymetry

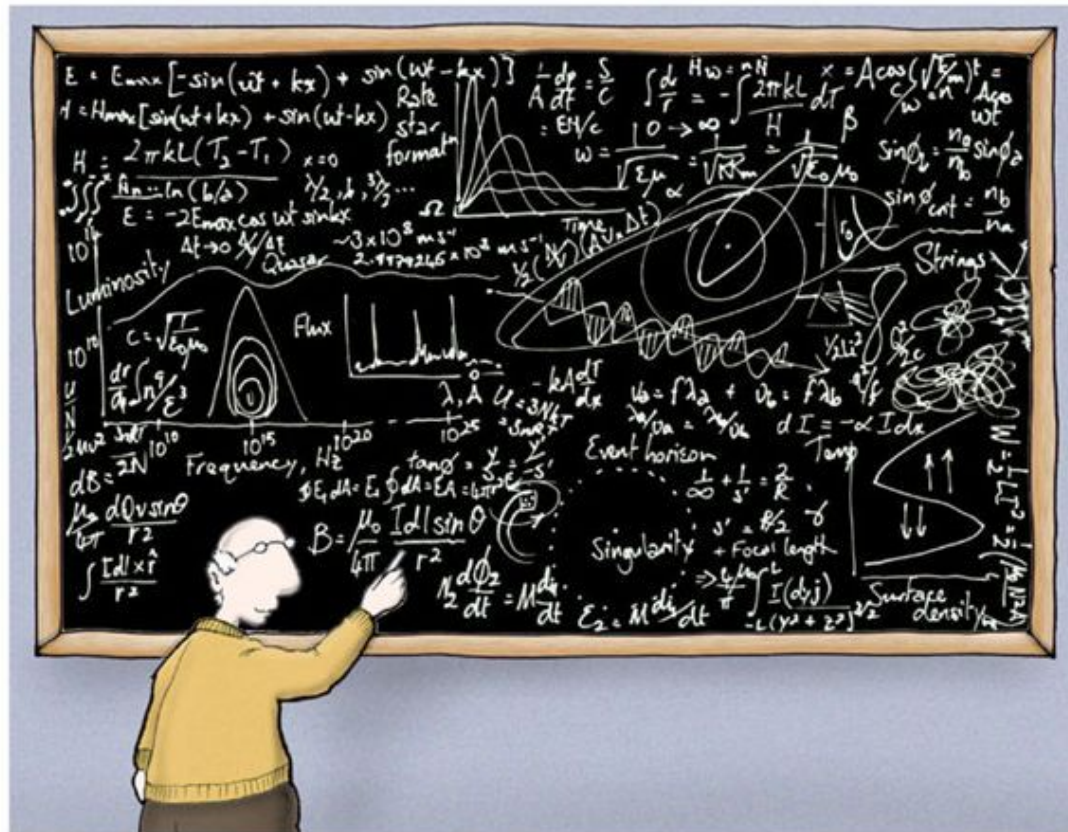
S

Simple questions: impact of the coastal shelf

- stabilize / destabilize ?
- unstable wavelength selection ?
- Size and vertical structure of detached eddies ?

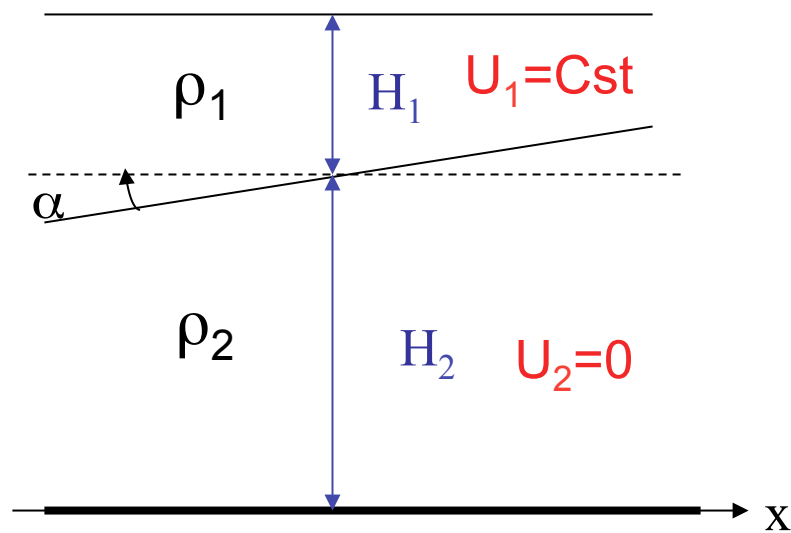
Surprisingly... no clear answers

What do we learn from linear stability analysis ?



What do we learn from simple models ?

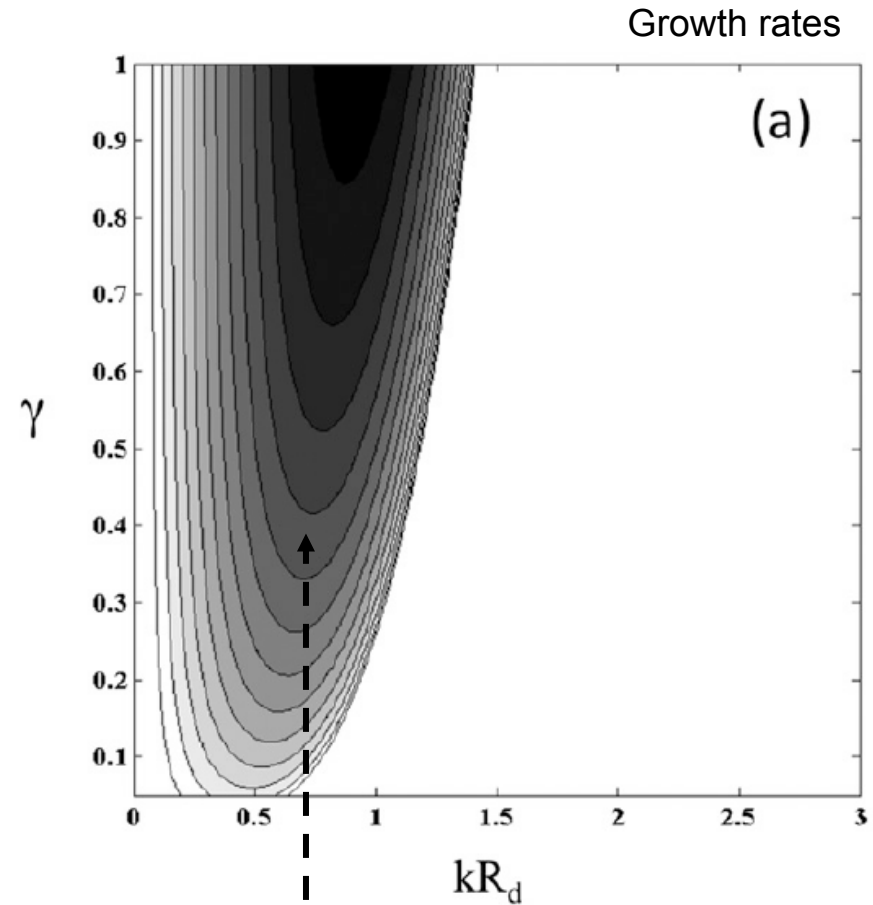
Phillips model, QG two layers



$$\gamma = H_1/H_2$$

$$h_2 \sim 2.5 h_1 \quad \gamma \sim 0.4$$

Bransfield current



$$kR_d \sim 0.65$$

$$\lambda_{\max} \sim 9 R_d$$

What do we learn from simple models ?

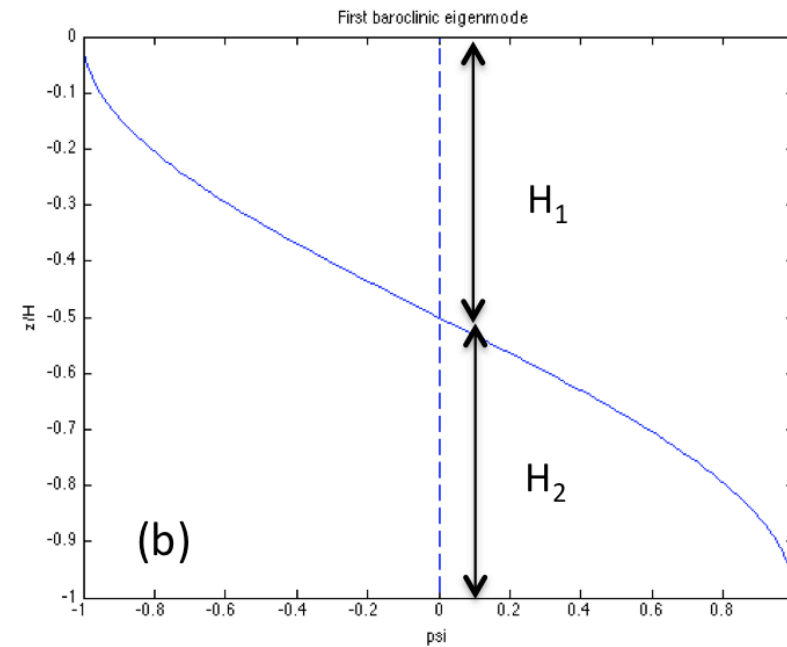
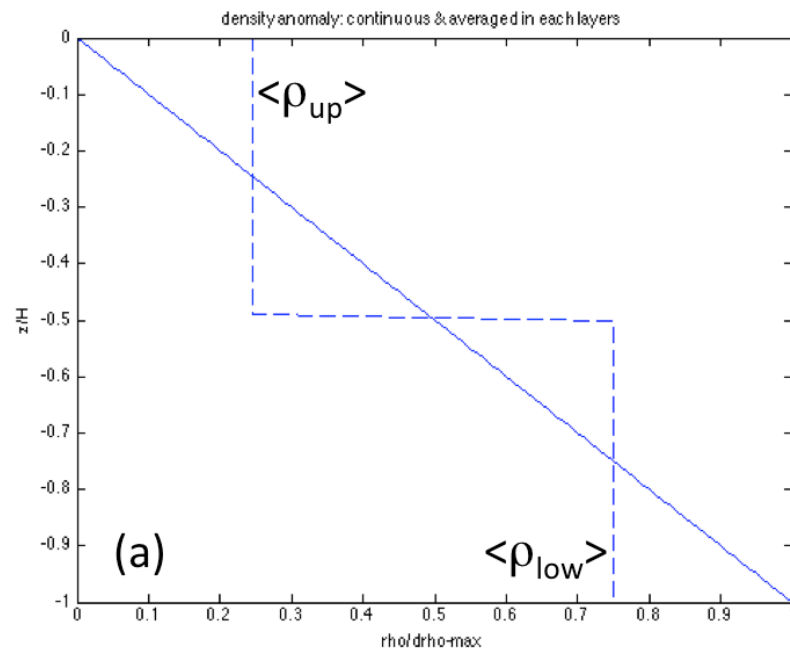
linear stratification

Stratification parameter

$$\gamma = H_1/H_2 = 1$$



$$N = Cst$$



What do we learn from simple models ?

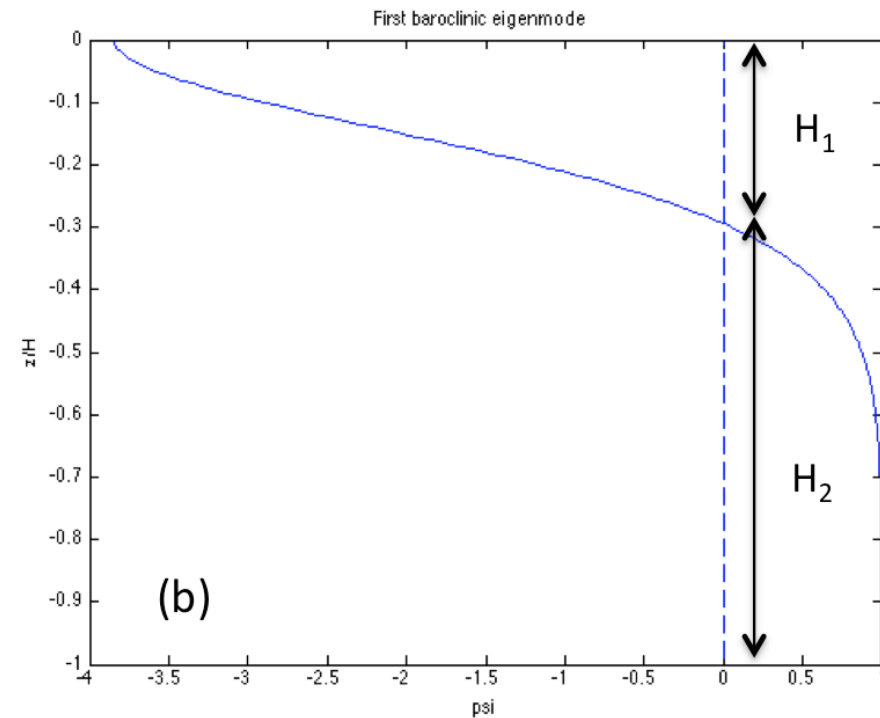
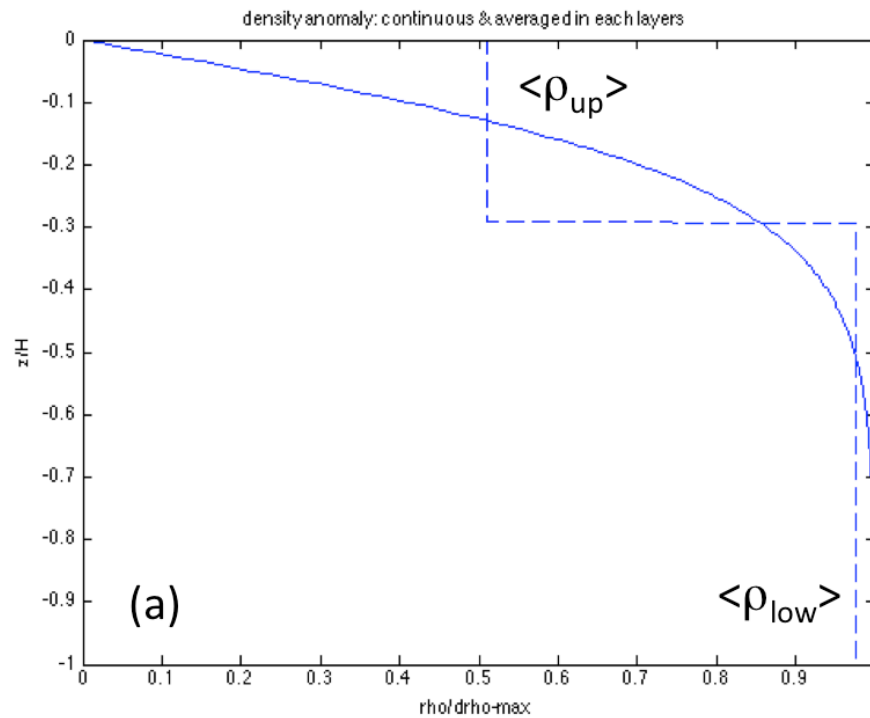
Stratification amplified at the surface

Stratification parameter

$$\gamma = H_1/H_2 < 1$$



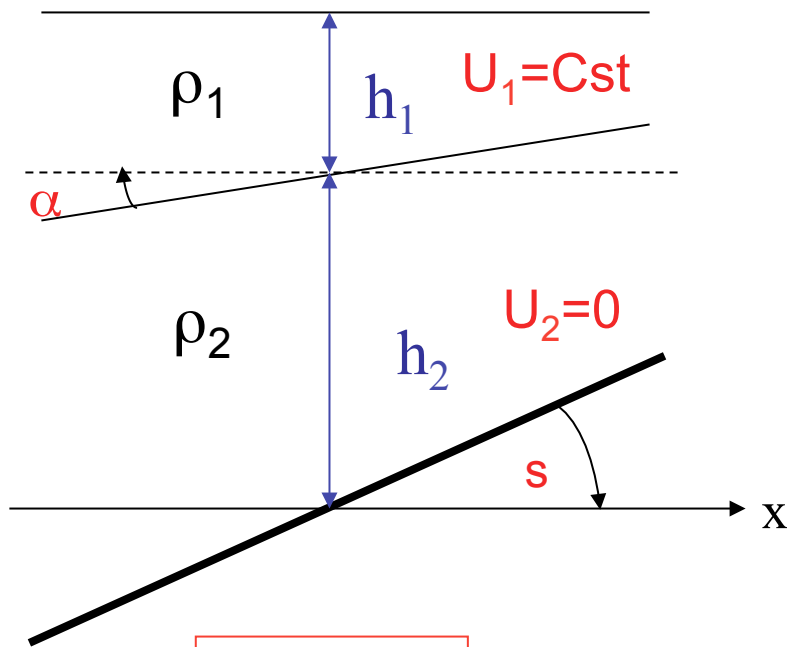
$$N \neq Cst$$



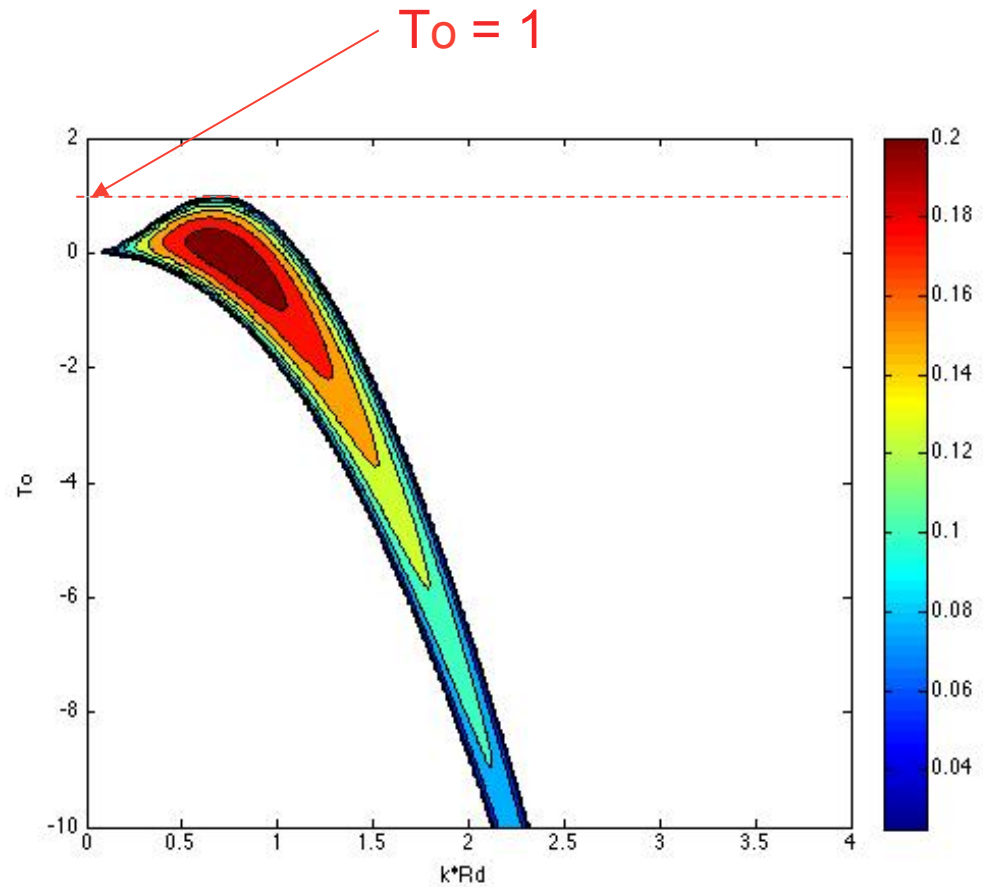
$$\gamma = 0.4$$

What do we learn from simple models ?

Phillips two layers QG model



$$T_0 = \frac{s}{\alpha} \geq 0$$



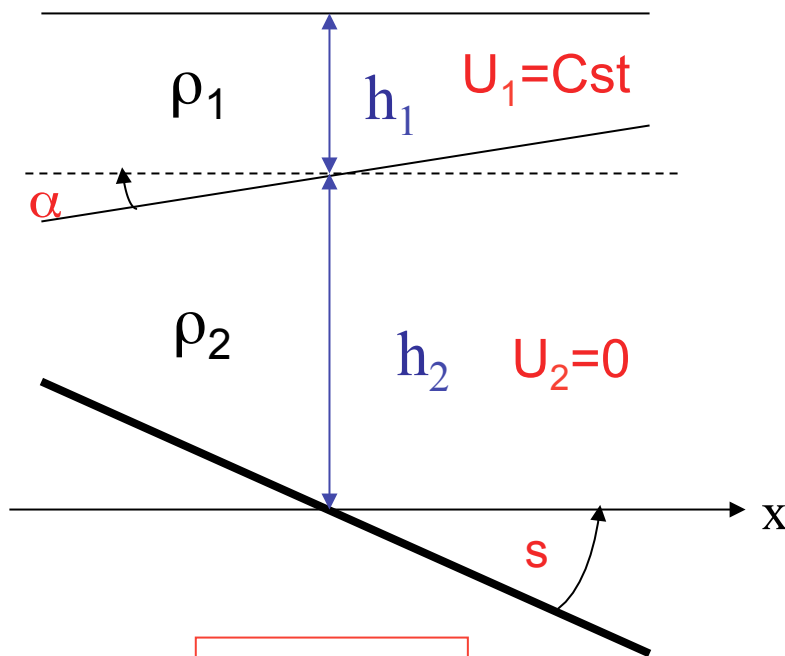
$T_0 > 1$



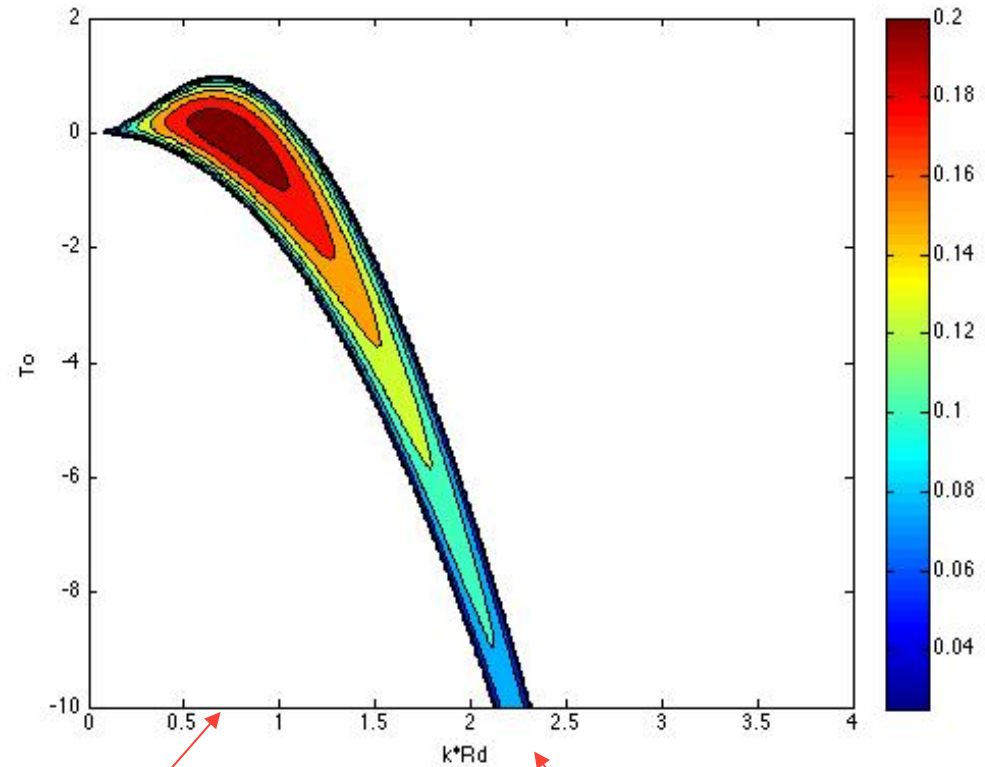
STABLE

What do we learn from simple models ?

Phillips two layers QG model



$$T_0 = \frac{s}{\alpha} \leq 0$$



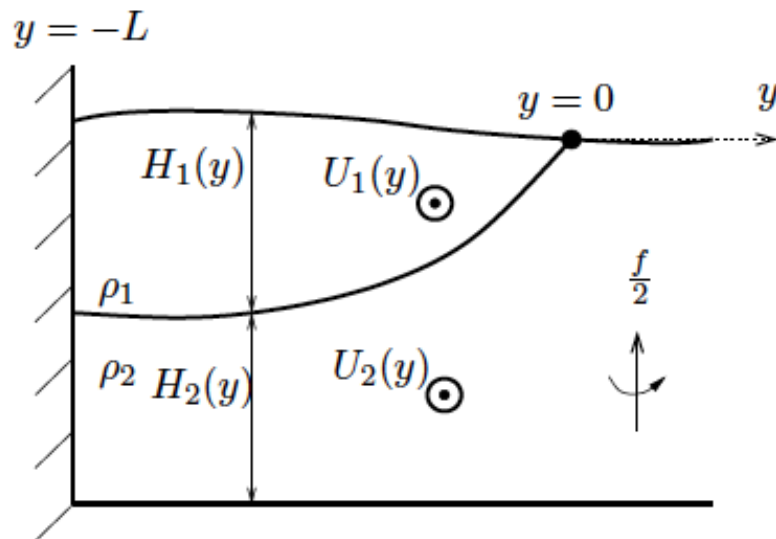
Long wavelength cutoff

Small wavelength selection & STABILIZATION

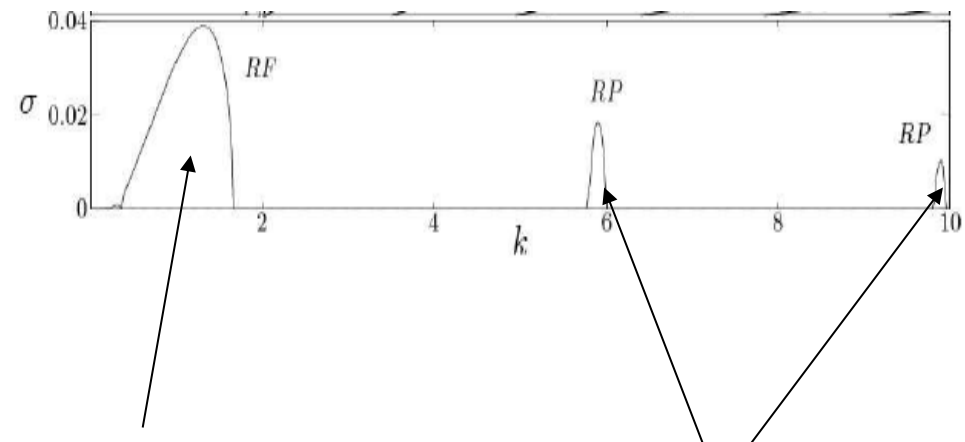
What do we learn from RSW models ?

Buoyant coastal front, flat bottom

Surface advected $\gamma = H_1/H_2 \sim 0.1$



$$L = R_d \quad Bu = 1$$



Frontal-Rossby modes
(i.e. Rossby-Rossby)

Poincaré-Rossby modes

Geostrophic instability

Ageostrophic instability

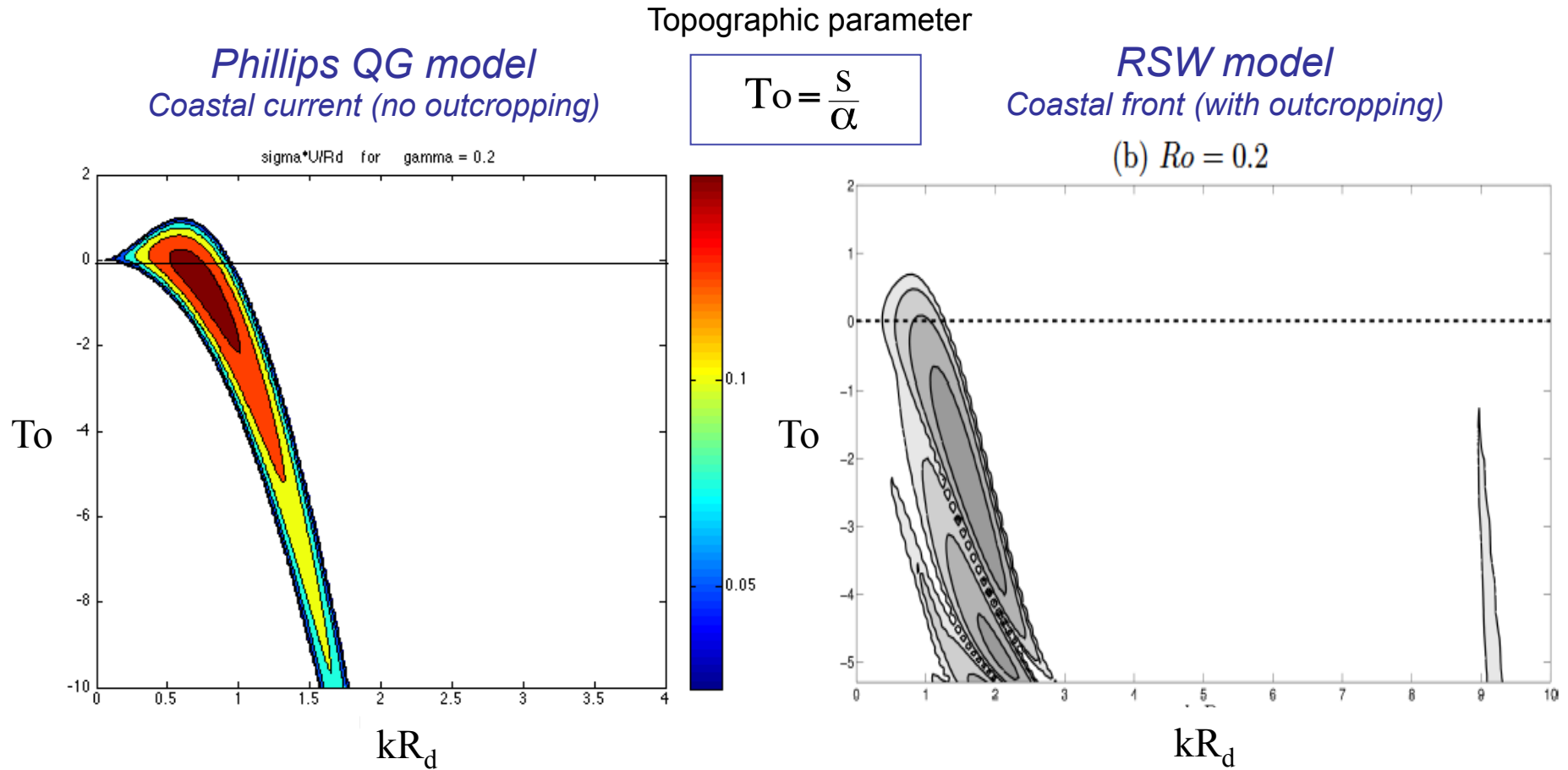
$$kR_d \sim 1.5$$

$$kR_d \sim 6$$

Boss, *et al.* (1996)

Gula, Zeitlin & Bouchut, *J. Fluid Mech.* (2010)

Impact of bottom slope (two-layer models)



Mysak JPO (1977)
Pennel PhD (2012)

$T_o < 0$



kR_d

σ_{max}

Gula & Zeitlin (2014)

Impact of bottom slope: **topographic parameter**

Upper layer *jet velocity*

$$V_{\text{current}} = \frac{g^*}{f} \partial_x h_1 = \frac{g^*}{f} \alpha$$

Phase speed of *Topographic Rossby Waves*

$$V_{\text{TRW}} \approx -s \frac{f}{Hk^2}$$

$$\frac{V_{\text{TRW}}}{V_{\text{current}}} = -\frac{s}{\alpha} \frac{f^2}{g^* Hk^2}$$

$$T_o = \frac{s}{\alpha} = -\frac{V_{\text{TRW}}}{V_{\text{current}}} k^2 R_d^2 \approx -\frac{V_{\text{TRW}}}{V_{\text{current}}}$$

when $kR_d \sim 1$

$$T_o < 0$$

$$V_{\text{jet}} V_{\text{TRW}} > 0$$

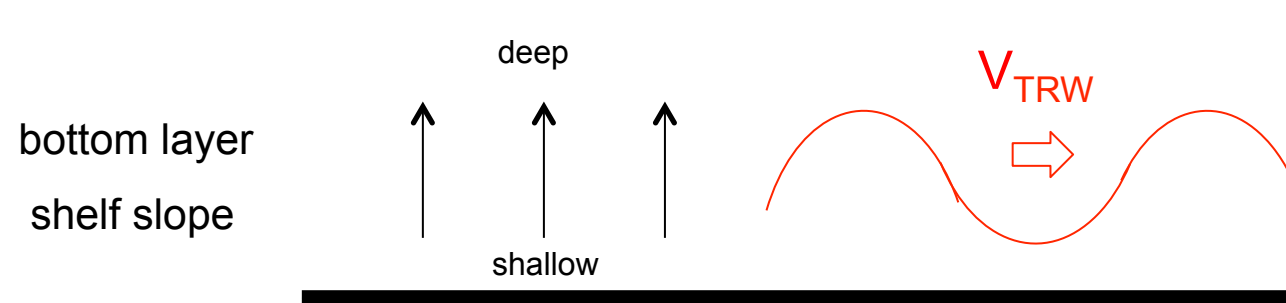
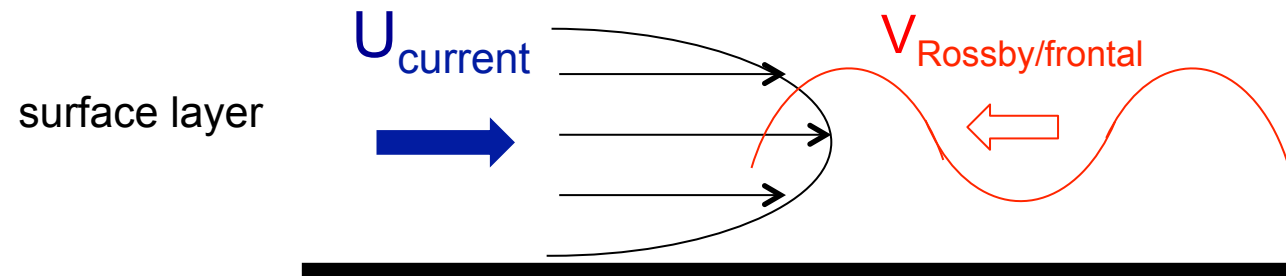
prograde topography

$$T_o > 0$$

$$V_{\text{jet}} V_{\text{TRW}} < 0$$

retrograde topography

Impact of bottom slope: simple mechanism

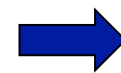


Topographic Rossby Waves

$$V_{TRW} \approx -S \frac{f}{Hk^2}$$

Phase speed locking

$$U_{current} - V_R = V_{TRW} + \gamma V_R$$



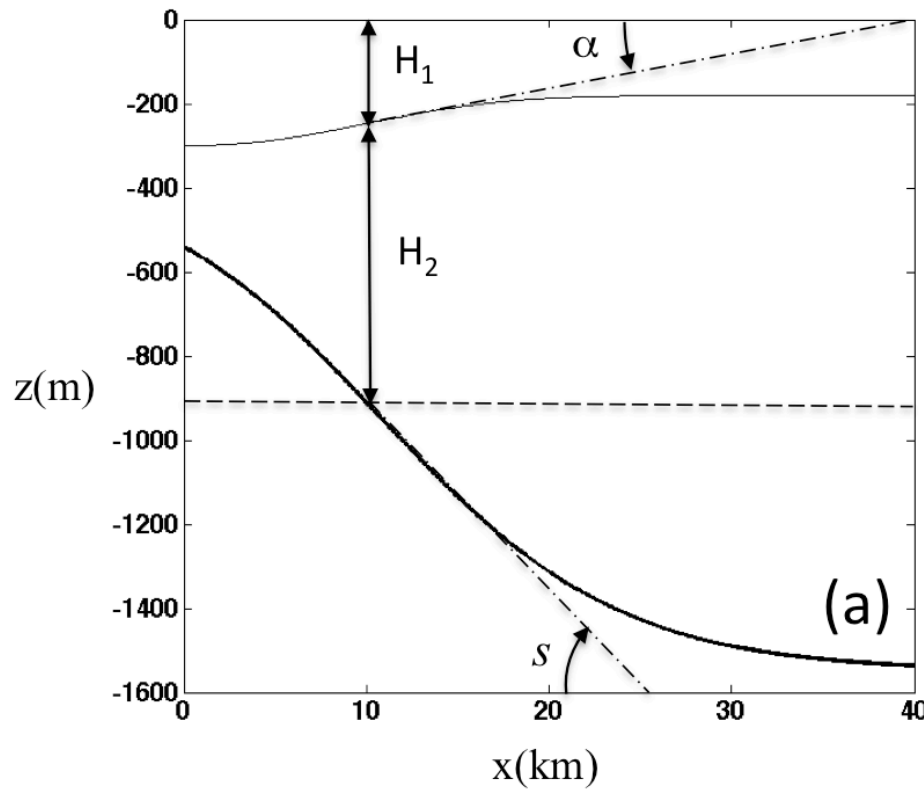
$$s < 0$$



$$kR_d$$

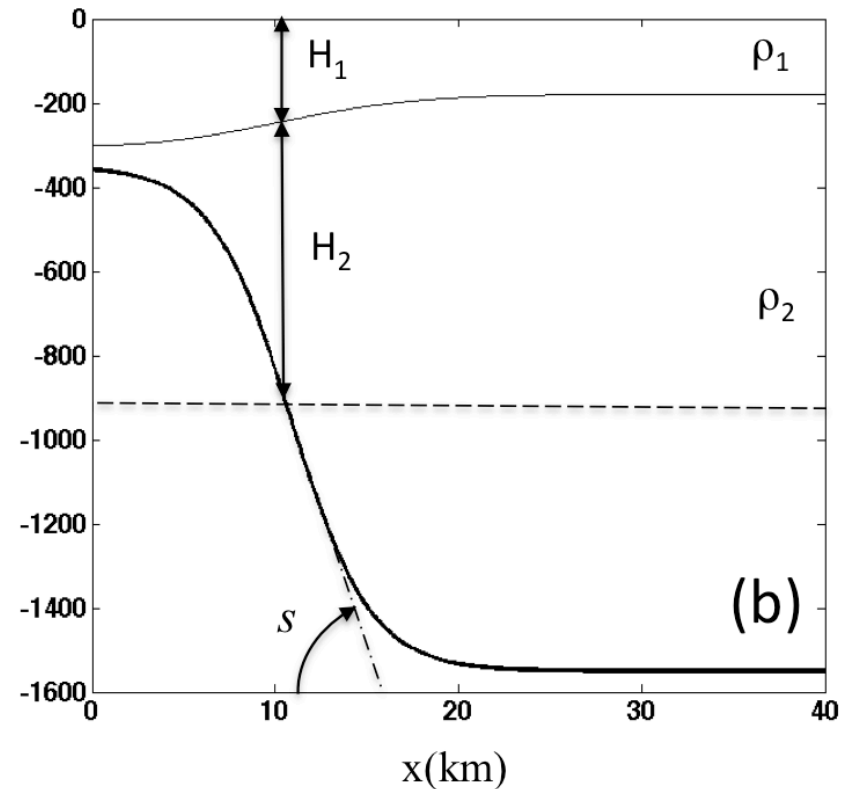


Shallow-water model: idealized two layers configuration



Vertical stratification parameter

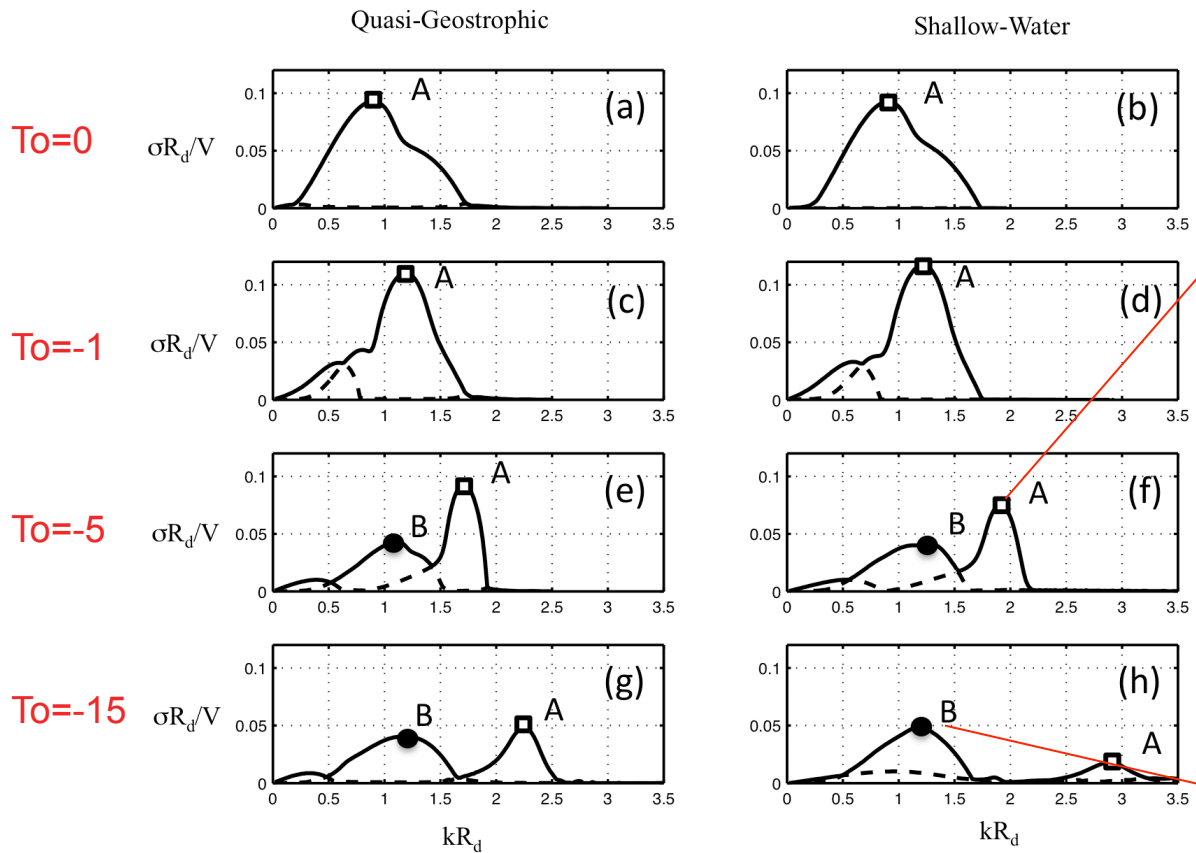
$$\gamma = H_1 / H_2 = 0.4$$



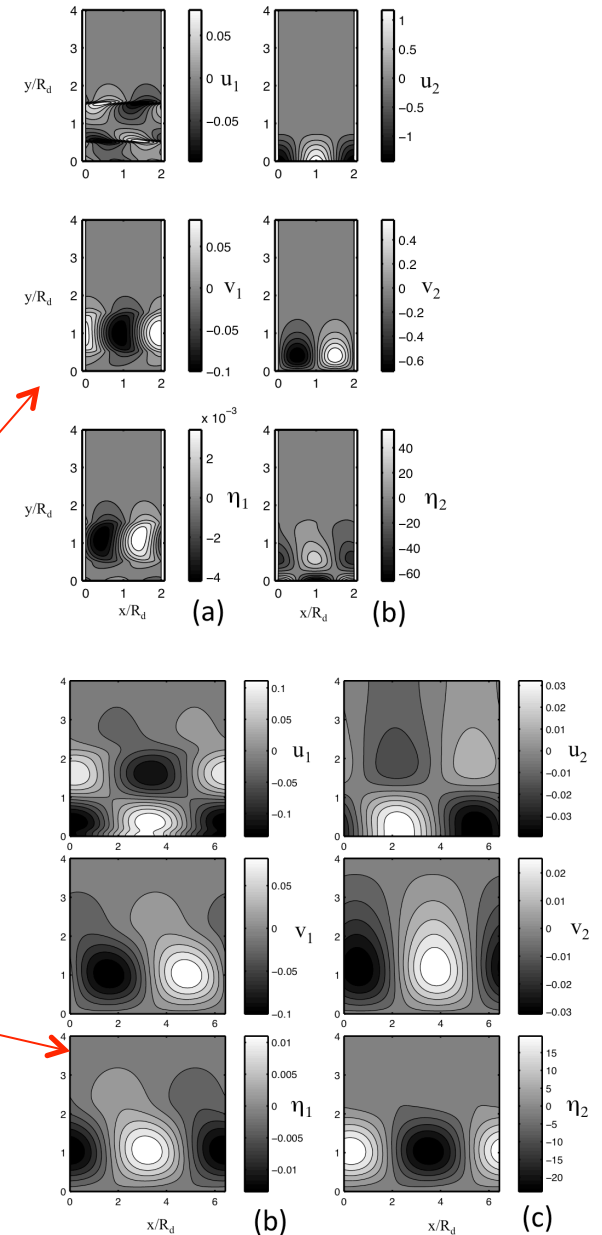
Topographic parameter

$$T_0 = s / \alpha < 0$$

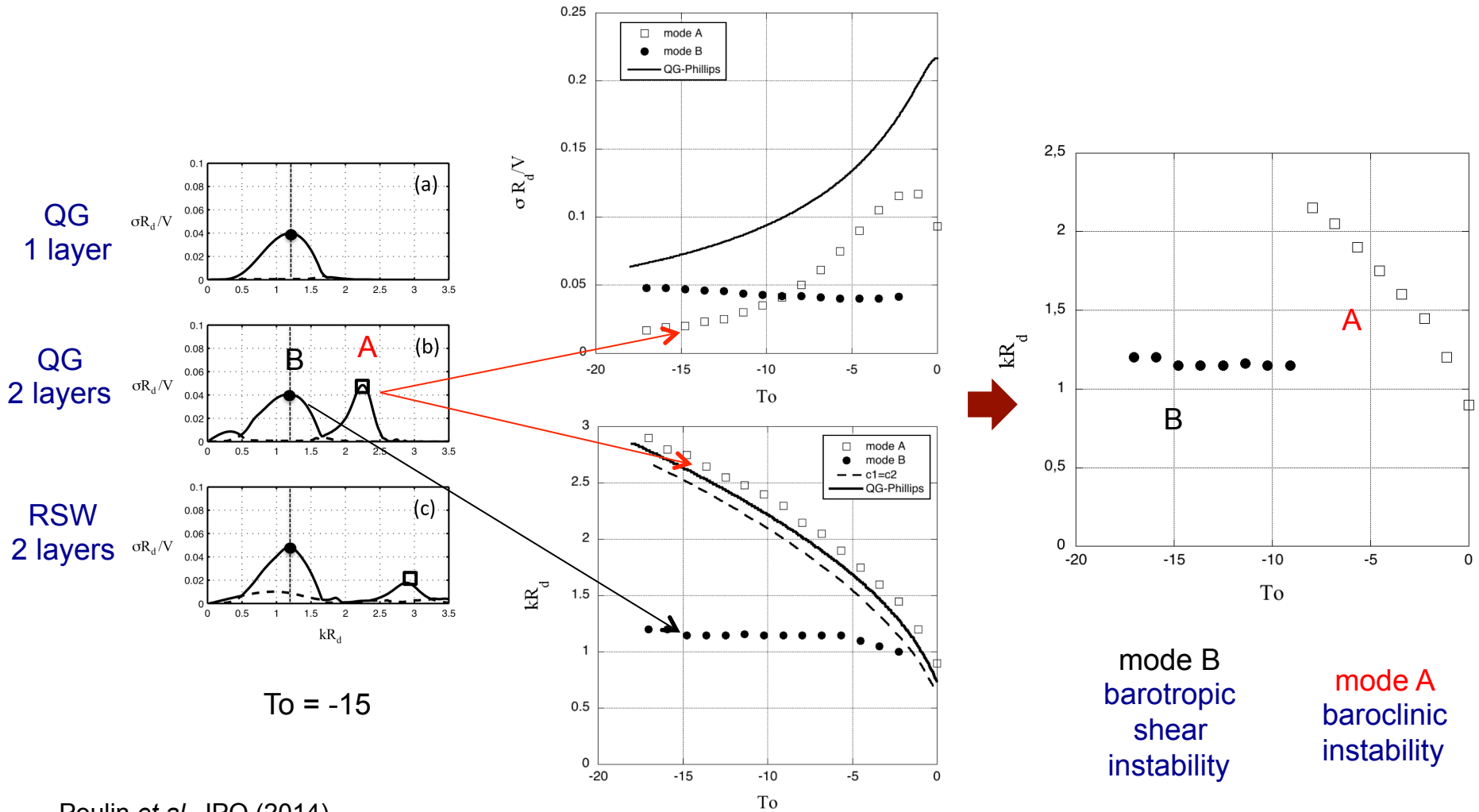
Shallow-water model: idealized two layers configuration



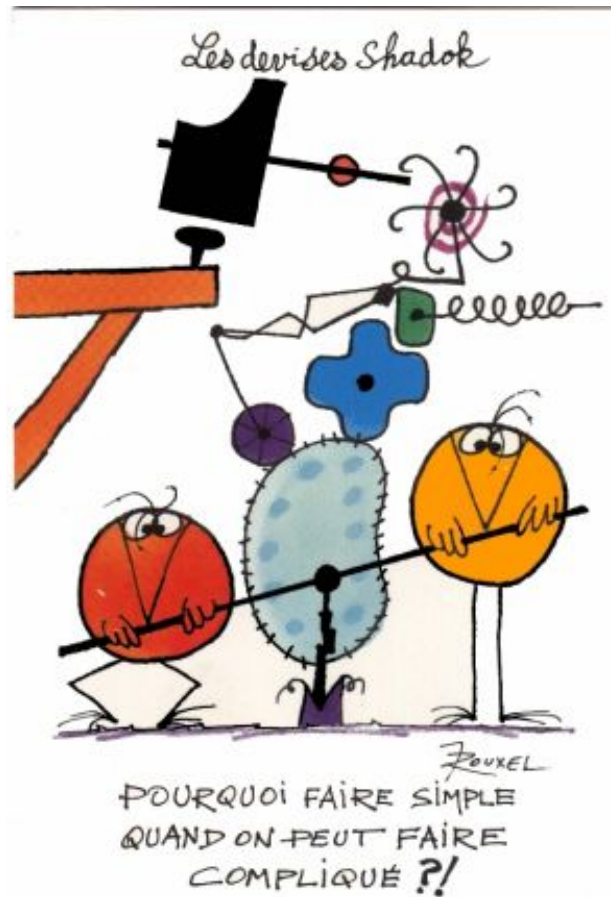
Two unstable modes: A,B



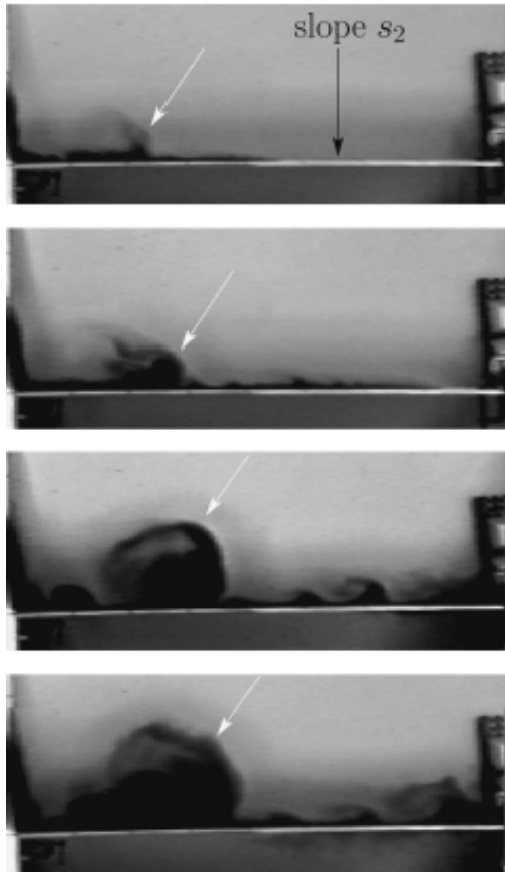
Shallow-water model: idealized two layers configuration



What do we learn from laboratory experiments ?

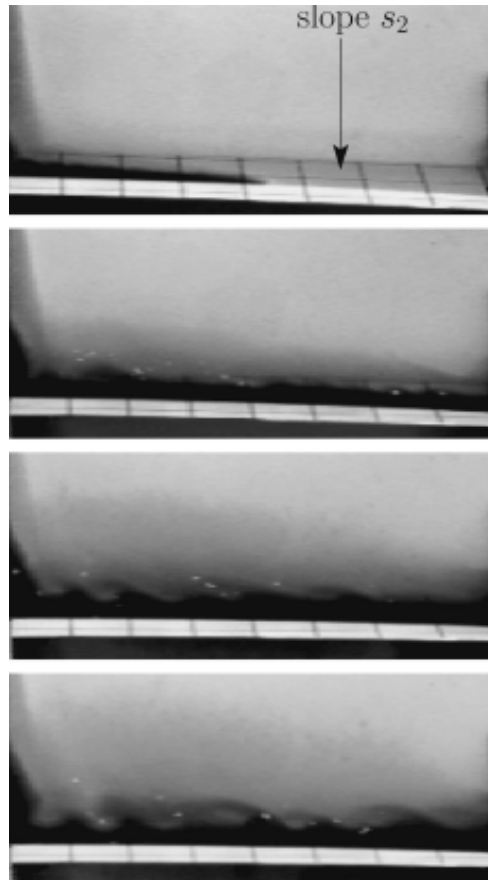


What do we learn from previous studies ?



$s_2 = 0$

formation of large eddy



$s_2 \sim 370\%$

smaller meanders

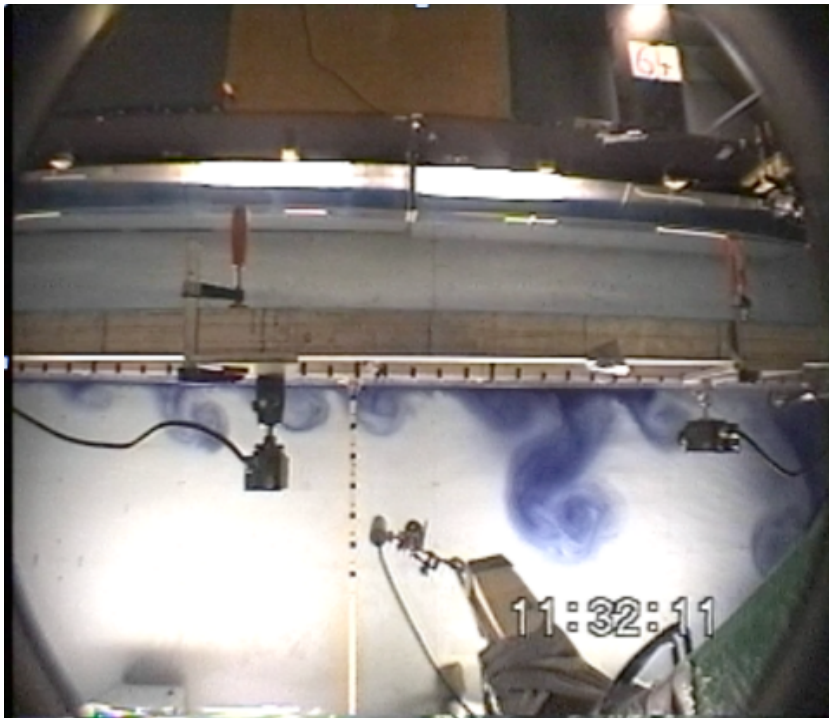
Surface advected configuration

$$H_1 \sim H_2/2$$

Wolf & Cenedese,
J. Phys. Oceanogr.(2006)

What do we learn from laboratory studies ?

Coastal gravity current in the Trondheim rotating platform (5m diameter)

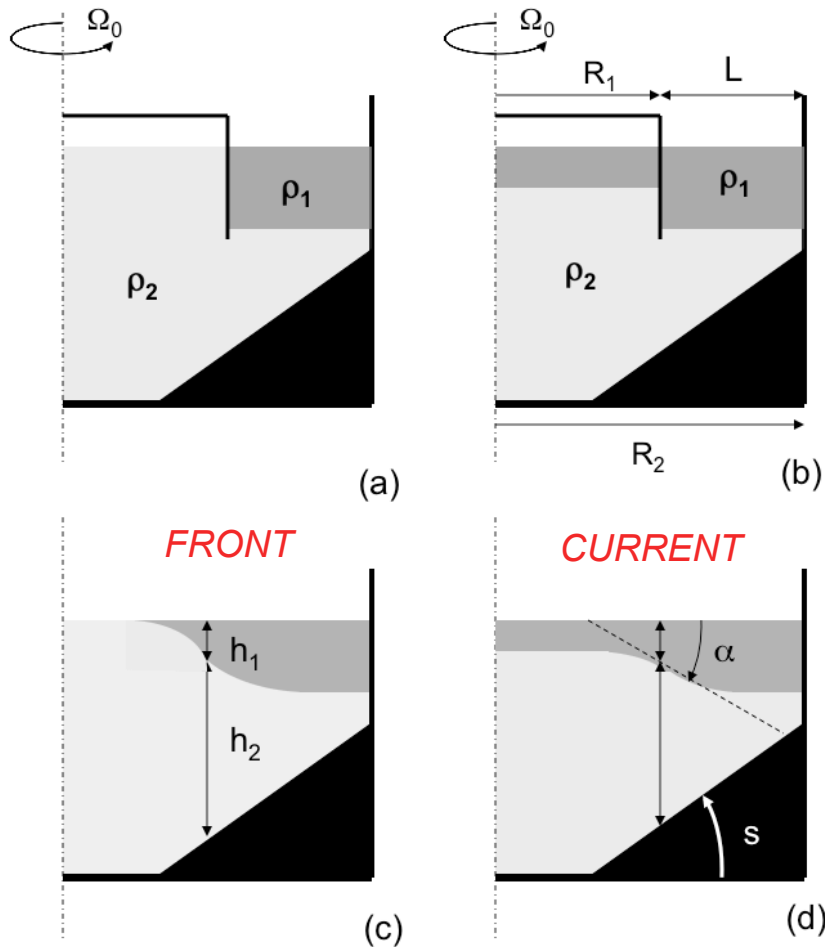


Flat bottom $s=0$

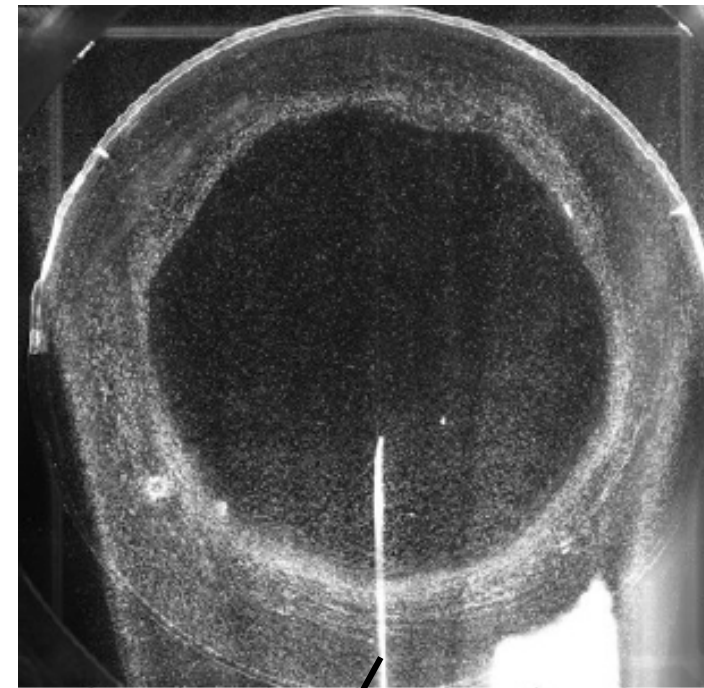


Steep slope $s=50\%$

Idealized configuration: experimental setup



*Initial and adjusted configurations
side view of the two layer salt stratifications*



top view with PIV particles



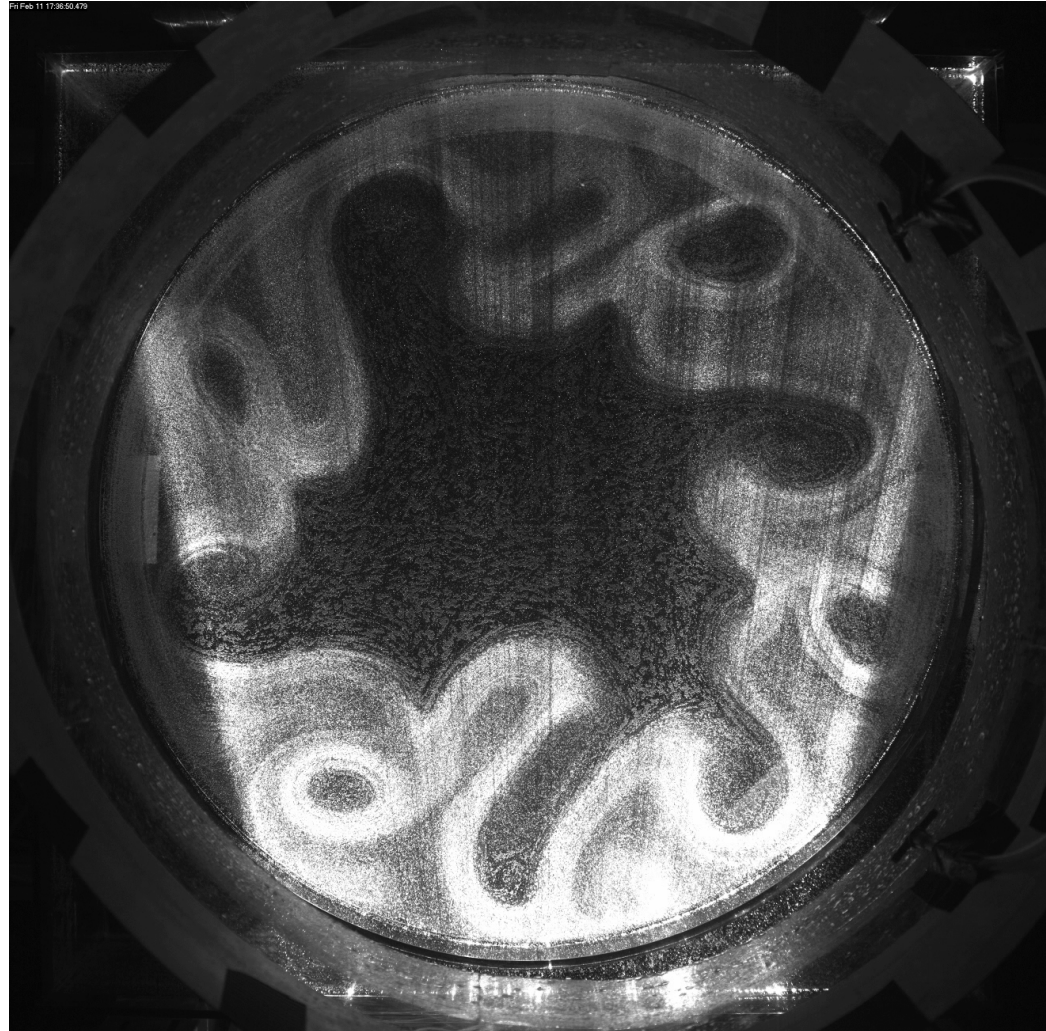
side view LIF visualization



Coastal current along steep shelf bathymetry



High resolution
PIV measurements:
4800x3200
pixels camera



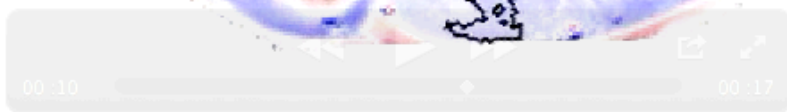
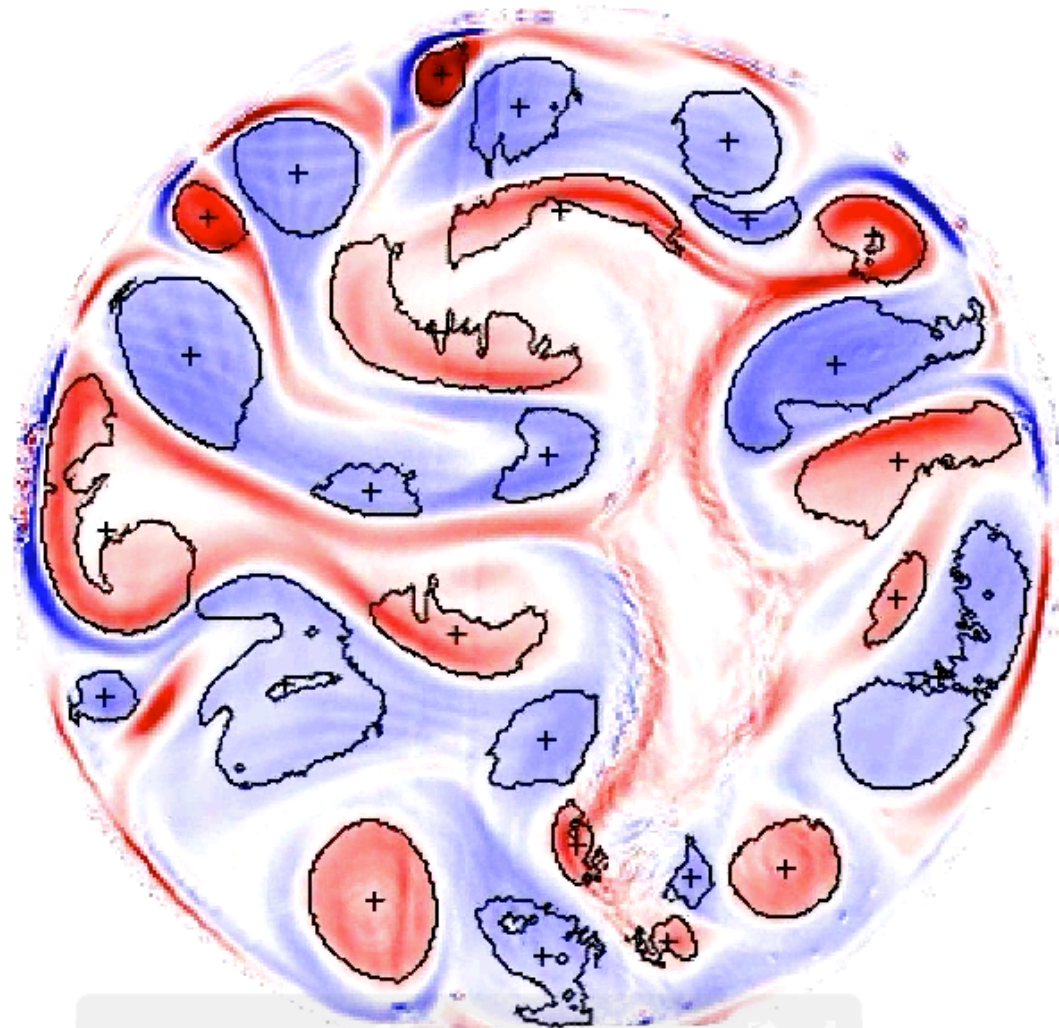
COASTAL FRONT
FLAT BOTTOM CASE

High resolution
PIV measurements:
surface vorticity

blue: anticyclonic
red: cyclonic

COASTAL FRONT
FLAT BOTTOM CASE

Black contours Okubo-Weiss criterion
(Isern-Fontanet et al. 2003)



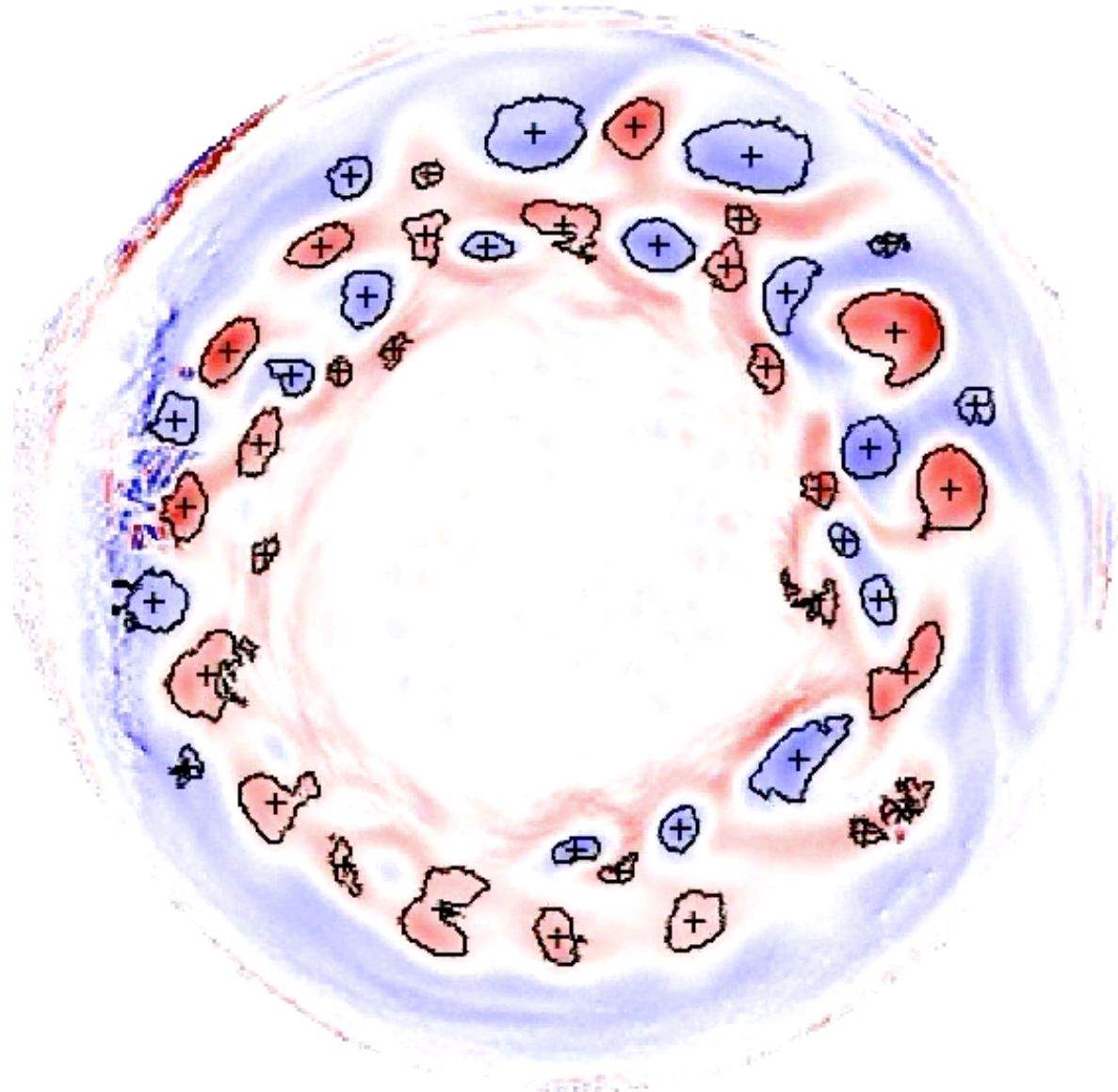
High resolution
PIV measurements:
surface vorticity

blue: anticyclonic
red: cyclonic

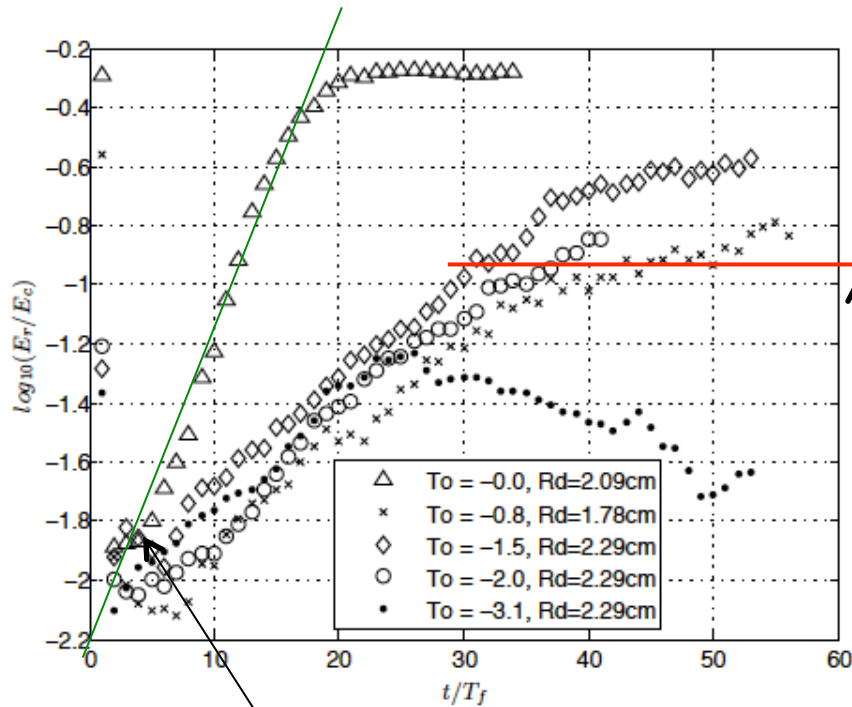
LINEAR SHELF

Topographic parameter

$$T_0 = \frac{s}{\alpha} = -1.3$$

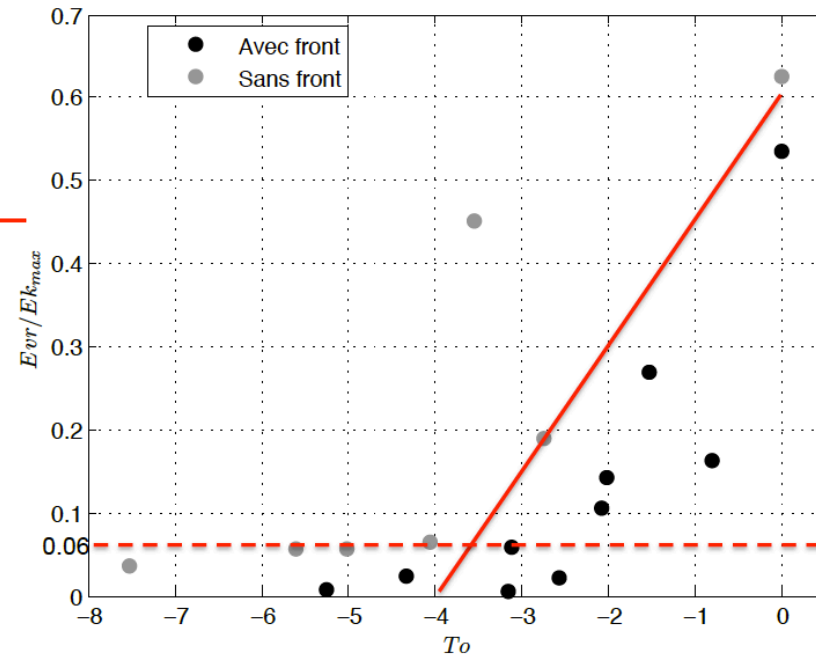


Evolution of cross shore KE/KE_{Tot}



Estimation of unstable growth rates
(flat bottom case $s=0$)

Level of non-linear saturation

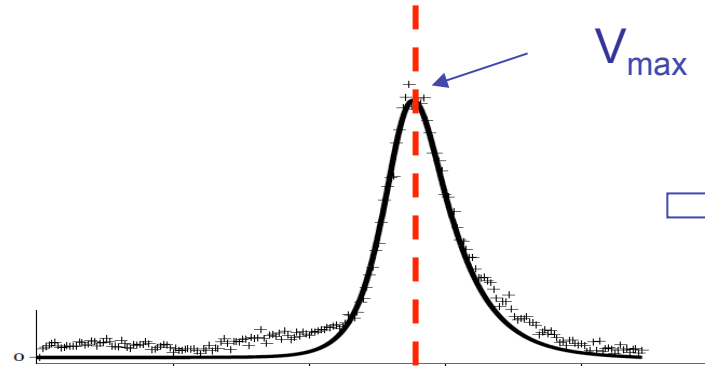


$$To_c \sim -3.5$$

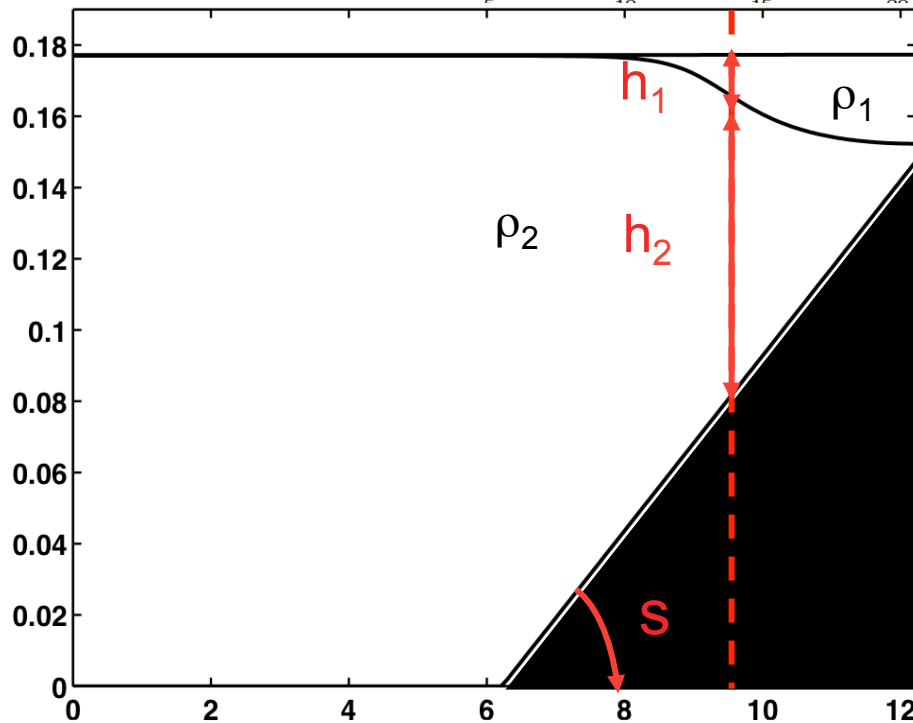
What do we learn from numerical simulations ?



Topographic impact: NEMO simulations / lab configuration



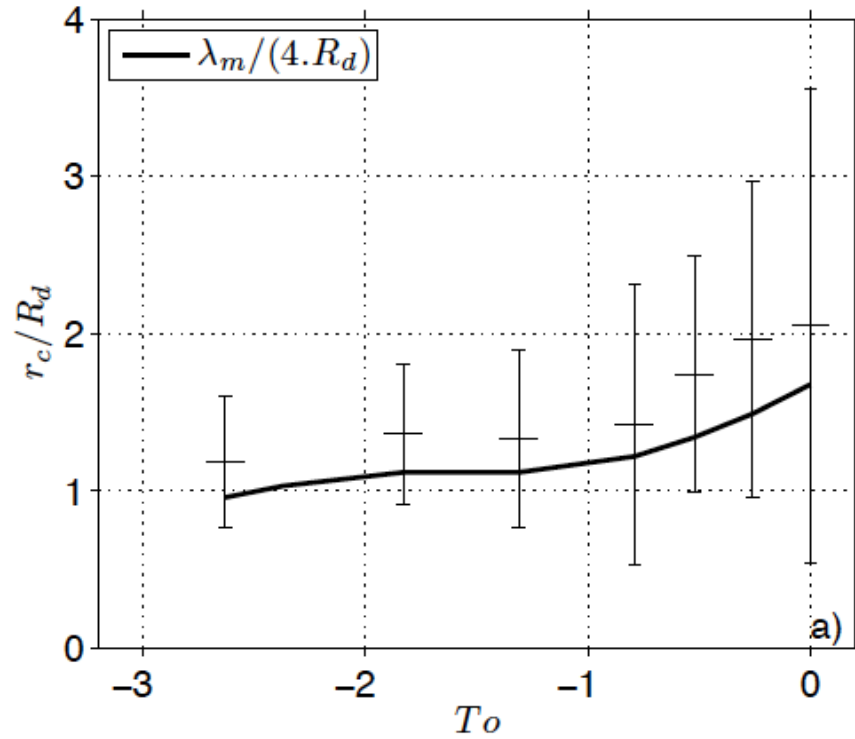
$$\Rightarrow Ro = V_{max} / (f R_d) = 0.35$$



$$\Rightarrow \gamma = h_1 / h_2 = 0.3$$

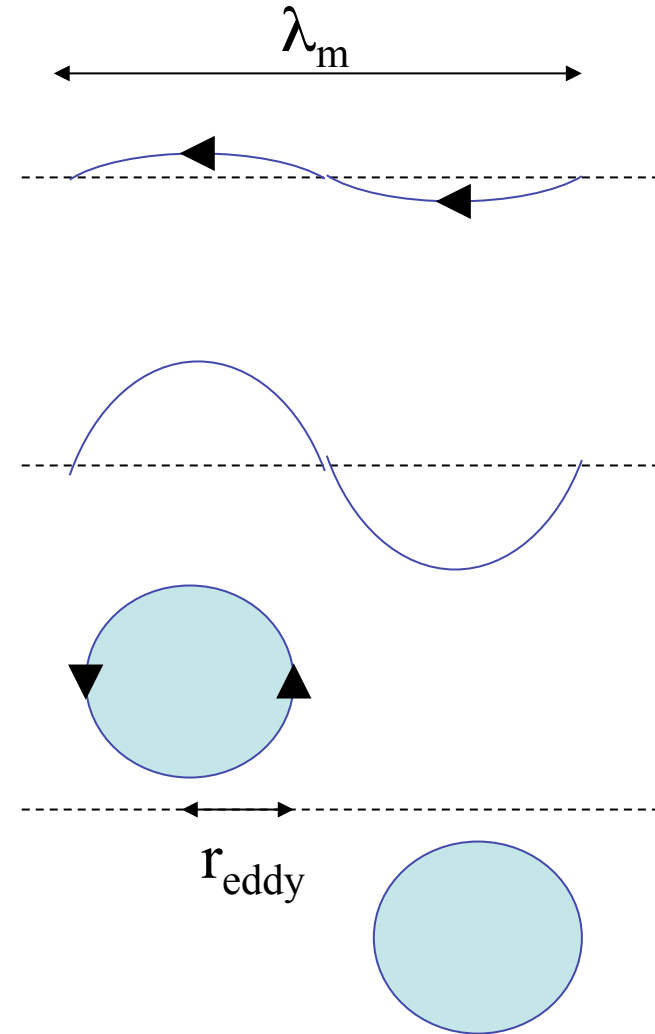
$$\Rightarrow To = s / \alpha$$

Topographic impact on: eddy size

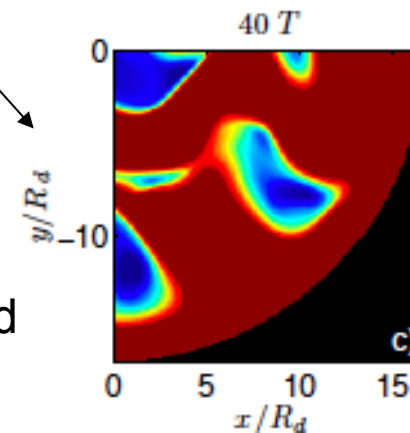
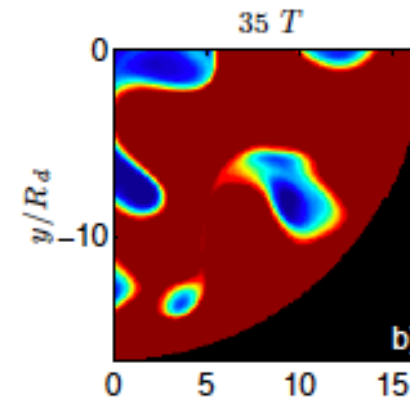
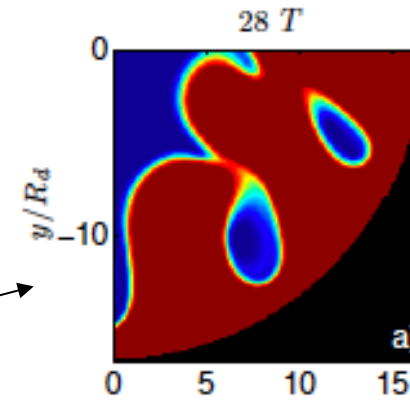
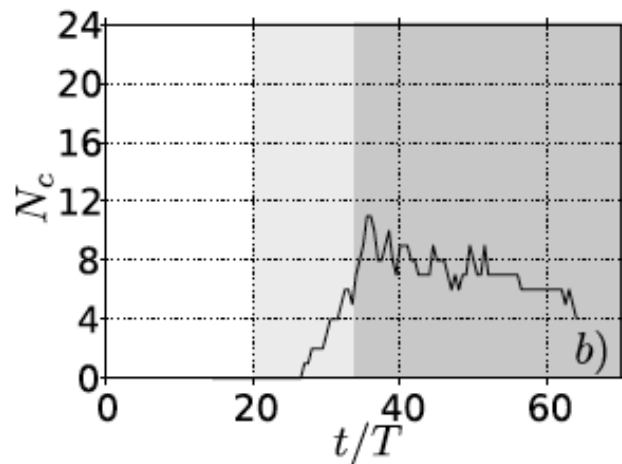
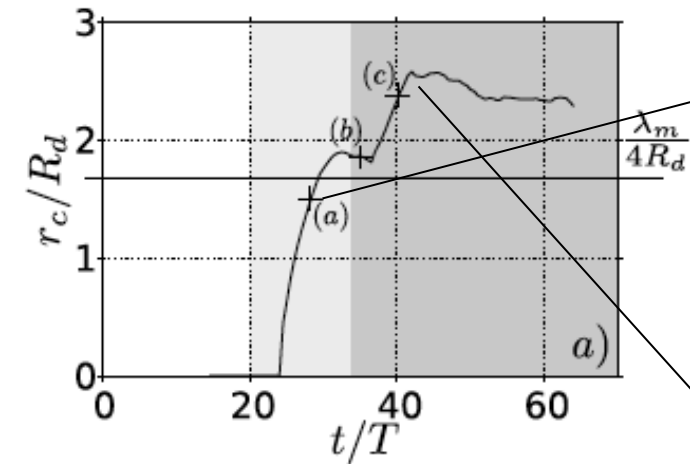


non-linear saturation

$$r_{\text{eddy}} \sim \lambda/4$$

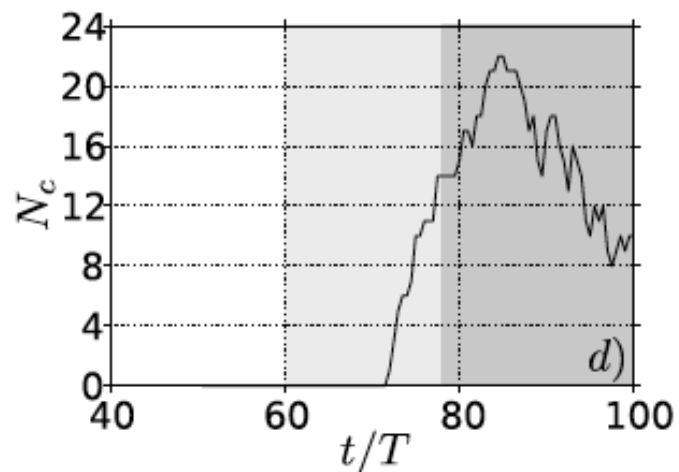
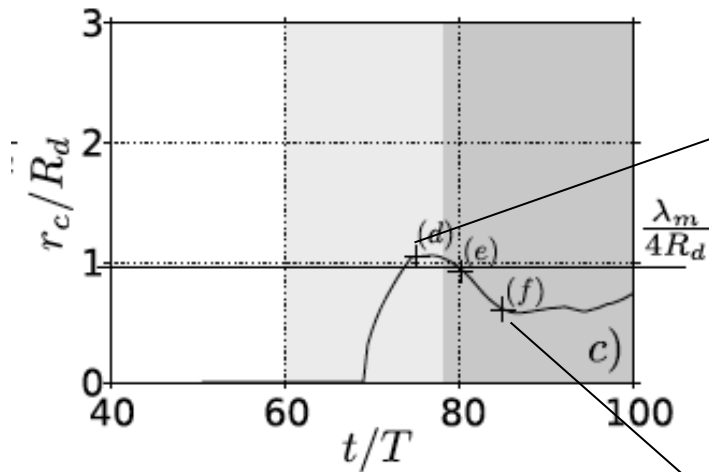


Secondary linear process: flat bottom $T_0=0$



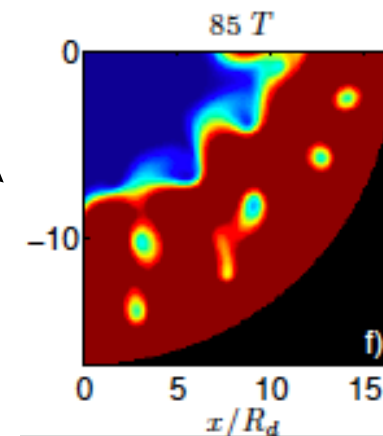
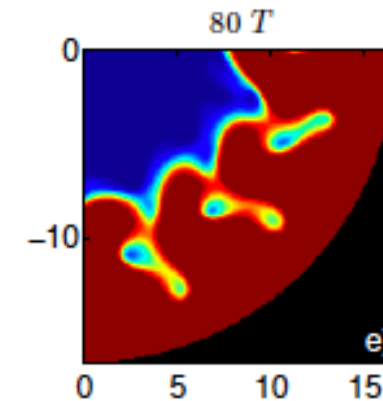
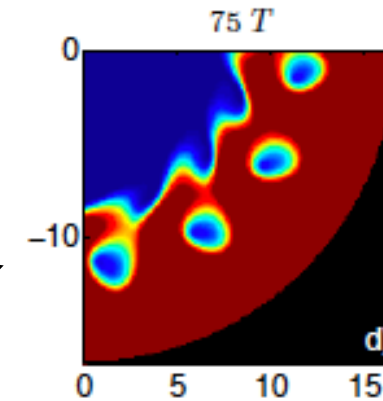
meso scale
cyclones
 $r_{\text{eddy}} \sim 2.5 R_d$

Secondary linear process: flat bottom $To = -2.6$



submeso scale
cyclones

$$r_{\text{eddy}} \sim R_d/2$$





CONCLUSIONS

Linear stability is not enough, non-linear shelf stabilization is crucial

Various dynamical regimes

- 1- Standard baroclinic instability, meso-scale anticyclones
strong cross-shelf exchanges
- 2- Rossby-TRW instability, **Trapped Coastal Instability**
reduced cross-shelf exchanges, small-scale trapped eddies

Direct **cascade towards sub meso-scale eddies** may occurs !

- 3- Strong non-linear stabilization, **weak** barotropic shear disturbances
no cross-shelf exchanges, strong along shore transport

Next step... next talk: fully stratified configuration (Bu , γ , T_p)